

Dynamics of PAHs and derived organic compounds in a soil-plant mesocosm spiked with 13C-phenanthrene

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| 1 | Dynamics of PAHs and derived organic compounds in a soil-plant |
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| 2 | mesocosm spiked with ¹³ C-phenanthrene. |
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14 Abstract

15 Polycyclic Aromatic Hydrocarbons (PAHs) are ubiquitous and persistent soil pollutants. 16 Their fate and the influence of the plant rhizosphere on their dynamics has been extensively 17 studied, but studies mainly focused on their dissipation rate. We conducted a plant-soil 18 mesocosm experiment to study the fate and distribution of PAHs or derived compounds in the 19 extractable fraction, the residual soil, the shoot biomass and the root biomass. The experiment 20 was conducted for 21 days using ryegrass and a forest soil spiked with ¹³C-labeled phenanthrene 21 (PHE), using combined IRMS and NanoSIMS for analyses. Almost 90% of the initial 22 extractable PHE content was dissipated within 3 weeks, but no rhizospheric effect was 23 highlighted on PHE dissipation. More than 40% of ¹³C-PHE was still in the soil at the end of 24 the experiment, but not as PHE or PAH-derived compounds. Therefore it was under the form 25 of new compounds (metabolites) and/or had been incorporated into the microbial biomass.

| 26 | About 0.36% of the initial ¹³ C-PHE was recovered in the root and shoot tissues, representing |
|----|---|
| 27 | similar ¹³ C enrichment (E ¹³ C) as in the soil (E ¹³ C \approx 0.04 at.%). Using NanoSIMS, ¹³ C was also |
| 28 | localized at the microscale in the roots and their close environment. Global ¹³ C enrichment |
| 29 | confirmed the results obtained by IRMS. Some hotspots of ¹³ C enrichment were found, with a |
| 30 | high ³² S/ ¹² C ¹⁴ N ratio. Comparing the ratios, sizes and shapes of these hotspots suggested that |
| 31 | they could be bacteria. |
| 32 | |
| 33 | Keywords: Polycyclic Aromatic Hydrocarbons; ¹³ C-Phenanthrene; soil; rhizosphere; IRMS; |
| 34 | NanoSIMS |
| 35 | |
| 36 | Highlights: |
| 37 | \triangleright A mesocosm experiment with ¹³ C-PHE-spiked soil was analyzed by IRMS and |
| 38 | NanoSIMS |
| 39 | > 40% of ¹³ C-PHE was still in the soil after 21 days but not as the parent compound |
| 40 | Bacterium-like ¹³ C-enriched hotspots were revealed by NanoSIMS close to the roots |
| 41 | |
| | |

42 **1. Introduction**

Polycyclic Aromatic Hydrocarbons (PAHs) are ubiquitous and persistent organic compounds found in the soils of large brownfield areas in many former industrial regions. They are composed of two or more benzene rings and result from the incomplete combustion by pyrolysis of organic materials from natural or anthropogenic sources (Edwards, 1983). They are persistent in soils due to their low water solubility and high adsorption on soil constituents (Wilson and Jones, 1993). They are harmful to human and environmental health, and 16 of them are listed as priority pollutants by the USEPA (Hartmann, 1996). 50 Many studies have addressed the fate and dynamics of PAHs in soils, where they can 51 undergo numerous processes such as plant uptake, stabilization by sorption on organic and 52 mineral phases, or transformation and degradation until total mineralization. Sorption and 53 degradation are key processes in the fate of PAHs in the environment (Magee et al., 1991). The 54 persistence of organic pollutants in soils has been attributed to sequestration mechanisms 55 including intra-organic matter and intra-particle diffusion of the parent compound and 56 transformation products (Doick et al., 2005). Plants seem to have a limited direct impact on 57 PAH dynamics in polluted soils (Fang et al., 2001; Günther et al., 1996; Rezek et al., 2008). 58 Gao and Zhu (2004) showed that uptake of phenanthrene and pyrene by plants accounted for 59 less than 0.01 and 0.23% of enhanced loss in vegetated versus non-vegetated soils. The low 60 water solubility of PAHs could prevent significant uptake by plant tissues (Binet et al., 2000; 61 Kipopoulou et al., 1999) and root-to-shoot translocation (Burken and Schnoor, 1998). However, 62 the indirect role of plants in PAH dynamics through their root systems is largely documented 63 (Rezek et al., 2008; Miya and Firestone, 2000; Liste and Alexander, 2000). The plant 64 rhizosphere is a favorable environment for high microbial diversity and activity, which permit 65 higher PAH dissipation in soils (Corgié et al., 2003; Anderson et al., 1993; Schwab, 1994; 66 Reilley et al. 1996). The main mechanisms involved in PAH dissipation in the soil and the 67 rhizosphere are biodegradation and biotransformation by an active microflora (Binet et al., 68 2000; Wilson and Jones 1993; Cerniglia, 1997). The soil heterotrophic microflora may use 69 organic pollutants as a source of carbon (C) and energy to grow, or for co-metabolism (Hwang 70 and Cutright, 2002). The C derived from PAHs (C-PAH) could be (i) incorporated into the 71 microbial biomass (Kästner et al., 1999), and/or (ii) transformed into other organic compounds 72 and thus form part of non-extractable PAHs (Richnow et al., 1998 and 1999). The newly formed 73 organic products could be more reactive than the parent compounds and therefore bind with 74 soil components more easily, but could also be more toxic (Sabaté et al., 2006). Cébron et al.

75 (2011) showed that 38% of the ¹³C from labeled PAH added into soil mesocosms was not 76 extractable after 12 days. C-PAH was still present but probably not under the form of parent 77 PAHs or extractable compounds. Some of these newly formed organic compounds may be more 78 soluble and more mobile than parent PAHs (Sabaté et al., 2006). However, most of the studies 79 led so far only considered the dissipation of PAHs in the soil and in the plant rhizosphere. Some 80 studies focused on the residual and newly formed organic compounds, especially on the 81 distribution and the dynamics of C-PAH in the soil (Doick et al., 2005; Eschenbach et al., 1998; 82 Kästner et al., 1999; Macleod and Semple, 2003; Northcott and Jones, 2001; Richnow et al., 83 1998), but no in soil-plant-microflora compartments.

84 An effective and relevant way to trace carbon compounds in soil-plant ecosystems is to use ¹³C isotopic labeling. For example, Richnow et al. (2000, 1998) used [1-¹³C]-PAH to trace 85 86 the transformation of phenanthrene in a soil reactor. But it appears more relevant to monitor C-87 PAH using a molecule labeled on all its carbon atoms (Cébron et al., 2011). Another 88 complementary method to trace a labeled organic compound in soils, plants or bacteria is to 89 map isotopic distribution at a microscale through NanoSIMS analyses (high-resolution imaging 90 Secondary Ion Mass Spectrometry) (Hoppe et al., 2013). NanoSIMS analyses can map all the 91 elements (isotopes) of the periodic table with an ultimate lateral resolution (probe size) in the 50 to 150 nm range, high sensivity (a few ppm), and high mass resolution (Herrmann et al., 92 93 2007; Moore et al., 2012; Remusat et al., 2012; Vidal et al., 2016).

The objective of our study was to investigate the fate of C-PAH in planted and non-planted soil by assessing (i) the distribution of C-PAH in the extractable fraction, the residual soil, the shoot biomass and the root biomass, and (ii) the localization of C-PAH in different compartments of the soil-plant system. We conducted a mesocosm trial using ryegrass, ¹³Cphenanthrene (¹³C-labeled on all its carbon atoms) soil spiking, and coupling isotopic analyses with NanoSIMS analyses.

100 **2. Materials & methods**

101 2.1. Soil-plant mesocosms

102 Planted and non-planted mesocosms were set up, with soil spiked with either unlabeled ¹²C-PHE (control) or labeled ¹³C-PHE. The soil was a Calcisol, sampled at 5-35 cm depth in a 103 104 temperate beech forest close to a former coking plant (49°14.163'N, 6°0.116'E, Homecourt, 105 Lorraine region, France). The soil horizon had a pH_{water} of 7.3; it contained 48% clay, 44% silt, 106 7% sand, had organic carbon and total nitrogen contents of 23.0 and 1.62 g kg⁻¹ respectively, 107 and a low PAH content (13.4 mg kg⁻¹). The soil sample was air-dried, ground and sieved to 2 108 mm, and then spiked with 300 mg PHE kg⁻¹ of soil. Spiking modalities were as follows: one 109 with unlabeled PHE (¹²C-PHE, Sigma Aldrich >97%) and two with known ¹³C-PHE/¹²C-PHE 110 mixtures (¹³C-PHE, Sigma-Aldrich Isotec Stables Isotopes, Miamisburg, Ohio, USA). The ¹³C-111 PHE/12C-PHE ratio was 1:8 for organic carbon analyses and 1:1 for NanoSIMS analyses. A 112 PHE-hexane solution was added to 1:10 of the total soil mass for spiking. Hexane was 113 evaporated under a fumehood, and then the spiked soil was homogenized with the remaining 114 9:10 of soil mass.

115 The planted and non-planted mesocosms were conducted in 50-mL glass jars with 40 g 116 of PHE-spiked soil. They were inoculated with 10 mL of soil suspension from a PAH-117 contaminated soil (200 g L⁻¹ in 0.85% NaCl) to ensure that a PAH-degrading microflora was 118 present. Five ryegrass (Lolium perenne) seeds were added in each planted mesocosm. The 119 experiment was conducted in a growth chamber under controlled conditions (22°C/18°C with 120 16h/8h day/night, 250 μ mol photons m⁻² s⁻¹, 80% relative humidity) for 21 days. Four replicates 121 were prepared for each planted and non-planted modality and each sampling time (0, 7, 14 and 122 21 days), and for each ¹²C-PHE- and ¹³C-PHE/¹²C-PHE- (1:8 ratio) spiked soil, except for 1:1 123 ¹³C-PHE/¹²C-PHE-spiked soil (one replicate, and sampling after 7 and 14 days for NanoSIMS 124 microprobe analyses). At each sampling time, ryegrass shoots and roots were removed from the soil and separated. The roots were washed with distilled water and dried with filter paper. The fresh plant material and soil were oven-dried at 35 °C and then ground, and the soil was sieved to 200 μ m prior to analyses. The roots were prepared directly after sampling for elemental mapping analyses.

129

130 2.2. Organic and isotopic analyses

PHE was extracted from 2 g of dry soil using dichloromethane (DCM) and a high-131 132 pressure and high-temperature DIONEX® 200 ASE (Accelerated Solvent Extractor) automated 133 extractor. Residual soil (after PAH extraction) was kept for organic analyses. DCM was 134 evaporated from the PAH extracts under N₂ flow and replaced by acetonitrile for PAH analysis. 135 PAH extracts were analyzed using a reverse-phase chromatography UHPLC Ultimate 3000 136 RSLC system (Dionex) equipped with a Diode Array Detector and a Zorbax Eclipse PAH 137 RRHD column (100 mm x 2.1 mm, 1.8 μ m – Agilent). The mobile phase was a mixture of 138 water/acetonitrile in gradient mode, with a flow rate of 0.420 mL min⁻¹. The wavelength used 139 for detection was 254 nm.

140

141 Isotopic and elemental analyses were performed on bulk and residual soil (before and 142 after PAH extraction, respectively), and on ryegrass shoot and root biomass using EA-143 GC/IRMS (Elementary analyzer - Isotope Ratio Mass Spectrometry 100, IsoPrime, 144 Manchester. EA: IsotopeCube, Elementar, Hanau). Approximately 1.0 mg of the powdered 145 material of each sample was weighed and placed in 6x4 mm tin capsules (Elemental 146 Microanalysis Limited, Devon, UK). Analyses were carried out at the Technical Platform of 147 Functional Ecology (OC 081) at the INRA Forest Ecology and Eco-Physiology Unit. Isotopic 148 measurements were reported in the delta notation (δ expressed in ‰) as:

149
$$\delta^{13}C(\%) = ((\frac{R_{max}}{R_{haf}}) - 1) \times 1000$$
 Eq. 1

150 Where $R_{sample} = {}^{13}C_{sample} / {}^{12}C_{sample}$ for labeled and control samples, and R_{Ref} refers to the Vienna 151 Pee Dee Belemnite (VPDB) standard, with $R_{Ref} = {}^{13}C_{PDB} / {}^{12}C_{PDB} = 0.0112375$.

152 Enrichment in ${}^{13}C$ (E ${}^{13}C$), expressed as excess % of atoms (at.%), was calculated relatively to

153 the unlabeled control in the soil and in the plant, as follows:

154
$$E^{13}C(at.\%) = ({}^{13}C_{labeled} - {}^{13}C_{control})$$
 Eq. 2

155 Where ${}^{13}C_{labeled}(at.\%) = [R_{labeled}/(R_{labeled}+1)] \times 100, {}^{13}C_{control}(at.\%) = [R_{control}/(R_{control}+1)] \times 100.$

156 E¹³C was calculated for bulk soil, residual soil, and ryegrass shoot biomass and root biomass.

157 Soil and plant ¹³C concentrations (¹³C_{conc}) were also calculated (expressed as μ g C g⁻¹ of soil and 158 μ g C g⁻¹ of plant, respectively):

159
$${}^{13}C_{ecose} (\mu g C g^{-1}) - (\frac{({}^{12}C_{black} - {}^{13}C_{black})}{100}) \times C_x$$
 Eq. 3

160 Where C_x refers to the C content, expressed in $\mu g C g^{-1}$ of soil or plant. ¹³C distribution (%) in 161 the mesocosms was deduced from the ¹³C concentrations and by considering that the percentage 162 of ¹³C initially added in the soil after PHE-spiking was 100% (T0). ¹³C distribution (%) was 163 calculated for the PAH extracts, the residual soil and the plants. The remaining ¹³C was 164 considered as ¹³C loss.

165 2.3. Elemental mapping analyses in the roots

The fresh ryegrass roots harvested after 7 and 14 days were rapidly cut into 0.5-cm pieces and placed on adhesive film. The fine roots were cryofixed using a LEICA EM CPC Universal Cryoworkstation. They were frozen at atmospheric pressure ($10^5 \, ^\circ C \, s^{-1}$), and slammed on copper mirrors (Printz et al., 2016). Then the cryofixed samples were lyophilized (LEICA EM CFD) without using any solvent ($1^\circ C \, h^{-1}$ for 8 days down to $-20 \, ^\circ C$ at 1.10^{-4} Pa) (Printz et al., 2016). The root samples were finally embedded in Epoxy resin and coated with gold prior to analyses by nano Secondary Ion Mass Spectrometry. 173 Elemental mapping was performed using a Cameca NanoSIMS 50 ion probe 174 (Gennevilliers, France) in the raster-imaging mode. The probe detects the secondary ions 175 generated from each pulverized sample surface by an energetic incident ion beam (Hoppe et 176 al., 2013). Ion images were obtained using a beam of Cs⁺ primary ions with an energy impact 177 of 16 keV and a beam current of 0.8 pA on the sample. The images were recorded in 256x256 178 pixels; under these conditions, a lateral resolution of 80-100 nm was expected. Negative 179 secondary ion signals for ¹²C⁻, ¹³C⁻, ¹²C¹⁴N⁻, ³²S⁻ ions and ¹²C¹⁴N⁻, ⁵⁶Fe¹⁶O⁻ clusters (Behrens et 180 al., 2008) were collected simultaneously using the multicollection system. Images were 181 acquired with a counting time of 30 ms/pixel. Mass resolution (M/ Δ M) was above 4500 (e.g. 2 182 916 was required to resolve mass interference between ¹³C and ¹²C¹H), and mass calibration 183 was achieved using standard iron oxide references. The isotope ratio of the two ¹²C and ¹³C 184 carbon isotopes allowed us to localize and quantify the ¹³C-labeled PHE. The quantitative data 185 acquired by NanoSIMS were calculated with Winimage software that directly processes image 186 data (CAMECA file format) and measures the $\delta^{0}/_{00}$ isotopic deviation according to the $\delta^{0}/_{00}$ 187 formula (Eq. 1) for each pixel. The deviation of the ${}^{13}C/{}^{12}C$ ratio was calculated based on a fixed 188 R_{ref} value (commonly 1.107%, the natural isotope ratio for ¹³C). The isotopic ratio of ¹³C 189 determined by NanoSIMS was not compare to IRMS values because of the presence of the 190 epoxy resin which prevent a specific ¹³C signal of organic matter (Vidal et al., 2016). The scale 191 of all the images of isotopic map deviation was kept constant (0 to 300) to allow for a quick 192 comparison of intensity levels.

Carbon-to-nitrogen (¹²C/¹²C¹⁴N) and sulfur-to-nitrogen (³²S/¹²C¹⁴N) ratio data were extracted from regions of interest (ROIs) by using the freeware package Image J. Each ROI was drawn at least twice for each group in order to obtain a mean ratio for any given area. Linescan (=segments of variable length drawn on the ion image) data were also extracted using OpenMIMS plugging with ImageJ software. Preparation and SIMS analyses were carried out 198 with a Cameca NanoSIMS 50 at the Luxembourg Institute of Science and Technology (LIST)199 in Esch-sur-Alzette, Luxembourg.

200

201 2.4. Statistical analyses

202 Statistical analyses were performed using R software (version 3.1.3, 2015). All tests were 203 used with a 95% significant level. Parametric conditions were checked with Shapiro-Wilk's 204 test for normality and with Bartlett's test for homogeneity of variance. When normality 205 conditions were not met, a non-parametric statistical test was used for multiple comparisons of 206 means. Two-way analysis of variance (ANOVA 2) followed by Tukey's post hoc test was used for the comparison of PHE and ¹³C concentrations between sampling times and between 207 208 treatments. For NanoSIMS analysis, the relative abundance levels of ¹³C from ROI were 209 compared with a non-parametric analysis of variance (Kruskal-Wallis test).

210

211 **3. Results**

3.1. Phenanthrene concentrations in planted and non-planted mesocosms

The extractable PHE concentration immediately after soil spiking (T0) was 171.6 ± 0.7 mg kg⁻¹ (Figure 1). After 7 days, PHE concentrations were 128.1 ± 9.7 and 137.6 ± 10.8 mg kg⁻¹ in the planted and non-planted mesocosms, respectively, indicating 20-25% PHE dissipation. The main decrease in PHE concentration was noted after 14 days, with 49.8 ± 15.6 and 40.2 ± 15.9 mg kg⁻¹ of PHE in the planted and non-planted mesocosms, respectively, i.e. approximately 70% of PHE loss. Finally, after 21 days there remained less than 20 mg kg⁻¹ of PHE in each mesocosm. The PHE concentration decreased significantly over time: 90% of the

added PHE was dissipated after 21 days. No significant difference in PHE concentrations was noted over time between the planted and non-planted mesocosms.

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3.2. ¹³C-enrichment and distribution in the planted and non-planted mesocosms

224 ¹³C-enrichment (E¹³C) values are summarized in Table 1 and Figure 2a. No significant 225 difference in E¹³C was noted between the planted and non-planted mesocosms. E¹³C of the bulk 226 soil decreased from 0.097 at.% at T0 down to 0.047 at.% after 21 days, while E¹³C of the 227 residual soil increased from 0.005 at.% to 0.044 at.% over the same time period. At the end of 228 the trial, there was no significant difference in ¹³C concentration between the bulk and residual 229 soils, indicating that ¹³C was mainly present in the residual soil under a non-extractable form 230 (Table 1). E¹³C in plant tissues was studied after 7 days. It reached 0.015 at.% and 0.017 at.% 231 in the roots and shoots, respectively. From 14 days, no significant difference was noted between 232 root and shoot E¹³C due to the high heterogeneity of values. After 21 days, E¹³C in residual soil 233 as well as in ryegrass shoots and roots was around 0.04 at.%.

234 ¹³C distribution was measured over time for the different compartments of the 235 mesocosms: residual soil, PAH extract, plant, and C loss (Figure 2b). At T0, 94% of the total ¹³C content was recovered in the PAH extract, and 6% in the residual soil. The ¹³C content of 236 237 the PAH extract decreased over time down to 3-11% of the total ¹³C content, while it increased 238 significantly in residual soil and reached 45-48% of the total ¹³C content after 21 days. The 239 ryegrass ¹³C content represented less than 1% of the total ¹³C content, mainly due to the very 240 low dry mass as compared to the total mass of soil (a few mg versus 40 g). The ¹³C loss from mesocosms accounted for 19% after 7 days and up to 43% after 14 days. 241

NanoSIMS analyses were carried out on unlabeled and labeled roots after 7 and 14 days. The mean percentages of ¹³C atoms in the roots were 1.09 in unlabeled (control) roots, 1.11 in ¹³C-labeled roots after 7 days, and 1.15 in ¹³C-labeled roots after 14 days. A significant ¹³Cenrichment was noted in labeled roots after 14 days relatively to the control.

248 A mapping of ¹³C associated with ¹²C, ¹²C¹⁴N, ³²S, ⁵⁶Fe¹⁶O ion species was performed on 249 root cross sections after 14 days (Figure 3). The roots were clearly identified in association with 250 high ¹²C⁻, ¹²C¹⁴N⁻ and ³²S⁻ signals. The resin was clearly visible with high ¹²C⁻ levels but no ¹²C¹⁴N⁻ or ³²S⁻ intensity, and was easily differentiated from the root tissue. Mineral particles 251 252 such as iron oxides, assimilated to soil particles, were highlighted by the ⁵⁶Fe¹⁶O - signal and 253 they were only observed surrounding the root. The composite image resulting from the overlay 254 of ¹²C, ¹²C¹⁴N and ¹³C elements emphasized ¹³C enrichment localization. There was no visual 255 evidence of global ¹³C enrichment in the roots, but rather locally distributed hotspots. Three 256 zones were identified close to the root epidermis (Figure 4), but none close to the soil particles. 257 These hotspots were mainly oval-shaped, and 1 to $2 \mu m$ in diameter.

258 An area of interest, pinpointed by the white square in Figure 3, was enlarged and 259 presented in Figure 4.a) Hotspots (red), root (green and blue association), soil (white) and resin 260 (green) were well identified. The black zone between the resin and the root could be holes that 261 formed during sample preparation. The hotspots were adherent to the root epidermal cells but 262 not to the soil particles. A linescan was drawn through root and hotspots (Figure 4.b) to study 263 ${}^{12}C^-$, ${}^{13}C^-$, ${}^{12}C^{14}N^-$ and ${}^{32}S^-$ distribution. It showed that a higher ${}^{13}C^-$ signal as compared to ${}^{12}C^-$ was 264 correlated with a higher ${}^{32}S^{-}$ signal as compared to ${}^{12}C^{14}N^{-}$. This higher ${}^{32}S^{-}$ signal as compared 265 to ${}^{12}C^{14}N^{-}$ was not observed in the root but only in the hotspot, so that the hotspots appeared to 266 be made of something else than root tissue. ¹²C/¹²C¹⁴N, ³²S/¹²C¹⁴N and ⁵⁶Fe¹⁶O/¹²C ratios were 267 also calculated for ROI selection in the hotspots, root tissue, soil particles and resin (Table 2).

Hotspots were easily discriminated from roots because they displayed significantly higher $^{12}C/^{12}C^{14}N$ and $^{32}S/^{12}C^{14}N$ ratios. They were not considered as soil particles because they had a low $^{56}Fe^{16}O/^{12}C$ ratio, or as resin because they contained less C than resin ($^{12}C/^{12}C^{14}N$ ratios equivalent to 12.88 for resin and 1.02 for hotspots).

272 **4. Discussion**

4.1. Effect of the rhizosphere on phenanthrene degradation

274 The extractable PHE concentration decreased by 40% after spiking and by up to 90%275 after 21 days in the planted and non-planted systems (Figure 1). This rapid dissipation of PHE 276 and its high rate are in part explained by experimental conditions, including the high 277 bioavailability of the freshly spiked pollutant and soil inoculation with adapted microflora. 278 Smith et al. (2011) showed that using a soil freshly spiked with pure PAH led to a greater loss 279 than using historically contaminated soils. Inoculation of a microflora extracted from PAH-280 contaminated soil had a priming effect in the first days following contamination and increased 281 PHE degradation (Joner et al., 2004, 2002). Low-molecular-weight PAHs such as PHE (3 rings) 282 can be used as primary substrates and be easily broken down in the soil by adapted microflora 283 commonly found in polluted soils (Cerniglia, 1997). This rapid PHE degradation could mask a 284 potential rhizospheric effect, even in a forest soil. No significant difference was noted in ¹³C-285 dynamics between the planted and non-planted systems (Figure 2). Therefore there is no 286 evidence of an impact of the rhizosphere on PAH degradation in this type of plant-soil 287 mesocosm.

288

4.2. Dynamics of phenanthrene carbon in the plant-soil mesocosms

290 The dynamics of PHE degradation was studied by monitoring the increase in C loss 291 and C incorporation in the soil (Figure 2.a). The loss of ¹³C during the first 14 days

represented half of the spiked ¹³C (Figure 2.b). We hypothesized that a major part of the
PHE loss was mineralized as CO₂. There was no significant difference in PHE degradation
between 14 and 21 days, which means that PHE degradation occurred during the first days
of contamination.

296 After 14 days, almost 40% of PHE carbon had been incorporated in the soil. This result 297 is in accordance with the 38% of 13C remaining in the soil after 12 days in comparable soil 298 mesocosms used to identify PAH-degrading bacteria (Cébron et al., 2011). No PHE metabolite, 299 like those identified by Käcker et al. (2002) and Richnow et al. (2000) for example, was 300 detected by GC-MS analysis of PAH extracts. The remaining 13C incorporated in the soil must 301 then have been present under the form of other organic compounds that were not extracted by 302 dichloromethane. Kästner et al. (1999) conducted a long-term experiment on the transformation 303 of 13C-labeled PAHs in soil bioreactors; they highlighted that the non-extractable residue had 304 a relatively high stability over time. These organic compounds, newly formed from PAHs, 305 become soil-bound residues (Binet et al., 2000) and/or a C source for microorganisms (Richnow 306 et al., 1998; 2000). For Richnow et al. (1999), the formation of these compounds most probably 307 depends on an intense microbial activity. Soil microorganisms may use organic pollutants like 308 PAHs as a C source and transform them into biomass through anabolic processes. Part of these 309 newly formed compounds can also be incorporated into organo-mineral complexes and be 310 protected against degradation depending on the nature of mineral phases (de Junet et al., 2013). 311 This could explain why these organic compounds remain non-available in the soil (Macleod 312 and Semple, 2003; Pignatello and Xing, 1996) and have a relative high stability (Richnow et 313 al., 1999).

Less than 1% of total PHE carbon was taken up by ryegrass roots and shoots. No significant difference was noted between the two plant parts after 14 and 21 days. Plant ¹³C enrichment was similar to soil ¹³C enrichment, namely about 0.04 at.% (Figure 2a), and

317 the difference expressed in total amounts was only due to the difference in mass between 318 the soil and the plants. However, according to NanoSIMS analyses (Figure 4), there was 319 no evidence of direct PHE uptake by plant roots. A small part of PHE may have been 320 sorbed onto the root epidermis, more precisely near the suberized exodermis and 321 endodermis (Dupuy et al., 2016). Previous studies show preferential PAH incorporation in 322 shoots through atmospheric deposition, and low direct incorporation by roots (Binet et al., 323 2000a; Desalme et al., 2011; Schwab and Banks, 1994). If ¹³C was not taken up by the 324 roots under the form of PHE, it may have been taken up under the form of newly formed 325 organic compounds or ¹³C-CO₂. It could be absorbed or adsorbed from ambient air by the 326 shoots after PHE volatilization (Gao and Zhu, 2004) or through stomata under the form of 327 ¹³C-CO₂ during the photosynthetic process. 13C enrichment was also noted in roots by SIMS, 328 confirming the excess ¹³C recovered in the roots (Figure 2.a).

329

4.3. Determination of ¹³C hotspots and localization of ¹³C in ryegrass roots

331 There was no correlation between the ⁵⁶Fe¹⁶O and ¹³C signals close to the roots after 14 332 days, therefore SIMS analyses did not evidence any contribution of organic matter-mineral 333 associations to the ¹³C-PAH dynamics. However two types of enrichment were highlighted 334 in the plant rhizosphere: a global ¹³C enrichment and ¹³C hotspots. Hotspots were composed of ¹²C⁻¹²C¹⁴N⁻ and ³²S⁻ ions species. Sulfur and nitrogen contents suggest the presence of 335 336 proteins (Eybe et al., 2008) that could originate from the plants or from microorganisms. 337 Hotspots were discriminated from root tissues based on their high ³²S content as compared to 338 the roots (Figure 4 and Table 2). The ${}^{32}S/{}^{12}C^{14}N$ ratio of the hotspots (0.11) was around five 339 times higher than the ${}^{32}S/{}^{12}C^{14}N$ ratio of the ryegrass roots (0.02). This last value is in accordance with the ³²S/¹²C¹⁴N ratio of 0.027 found by Dijkshoorn and Van Wijk (1967) in 340 341 ryegrass shoots. The oval shape of the hotspots and their size (1 to 2 μ m in diameter) are in accordance with the shape and size of bacteria. Moreover, measurements were performed on
isolated bacteria (unpublished data) and showed similar ratios (0.1 for ³²S/¹²C¹⁴N). There was
evidence of global ¹³C enrichment during plant growth, but no PAH uptake by the roots was
visualized by NanoSIMS analyses (Figure 4). The main ¹³C-PAH uptake could be attributed to
the bacteria present in the rhizospheric environment.

347

348 **5.** Conclusion

349 A mesocosm experiment was conducted using ¹³C-phenanthrene to monitor the 350 dynamics of PHE and derived compounds in the soil and in ryegrass. After 21 days, almost 351 90% of spiked PHE had been degraded. PHE was mainly mineralized under gas phase (up to 352 53%, corresponding to the ¹³C loss). But after 14 days, more than 40% of PHE carbon had been 353 transformed into newly formed organic compounds (metabolites) or incorporated into the 354 microbial biomass. No rhizospheric effect on PAH dynamics was noted due to the rapid 355 dissipation of PHE. Less than 1% of PHE carbon was also found in the root and shoot tissues 356 of ryegrass. The plant $E^{13}C$ was similar to the soil $E^{13}C$ (~0.04 at.%), suggesting a homogeneous 357 C-PAH distribution in the plant-soil mesocosm.

358 Global root C-PAH enrichment was also determined in the ryegrass rhizosphere at a 359 finer scale using NanoSIMS, based on ¹³C/¹²C ratio measurements on images obtained for 360 control (12C) and labeled samples after 7 and 14 days. There was no evidence of ¹³C-PAH 361 association to iron oxides in the close root environment. However, hotspots of ¹³C enrichment 362 close to the roots probably indicated the presence of bacteria, as suggested by their size, shape 363 and element composition. A part of PAH carbon was sorbed on the roots, but it was mainly 364 incorporated into the bacteria mentioned above. These bacteria probably degraded PHE and 365 incorporated PHE-derived ¹³C into their biomass.

The use of labeled contaminants and the coupling of IRMS and NanoSIMS is a relevant 366 367 and powerful tool for determining the dynamics of these contaminants. The dynamics of ¹³C 368 from labeled PAH was highlighted by ¹³C-partionning at a global level, but the transformation 369 of PAH ¹³C at a molecular level remains to be investigated. Complementary techniques such as 370 ¹³C-NMR and pyGCMS could be used to further characterize the ¹³C-PAH metabolites. 371 372 Acknowledgements 373 This work was supported by a grant from the University of Lorraine and the Lorraine 374 Region. The authors thank David Billet for UHPLC analyses and Patrick Grysan for NanoSIMS 375 analyses. 376 References 377 378 Anderson, T.A., Guthrie, E., Walton, B.T., 1993. Bioremediation in the rhizosphere. Environ.

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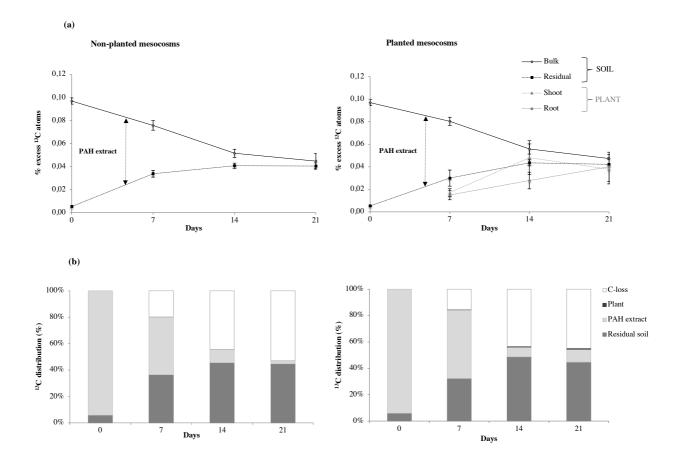
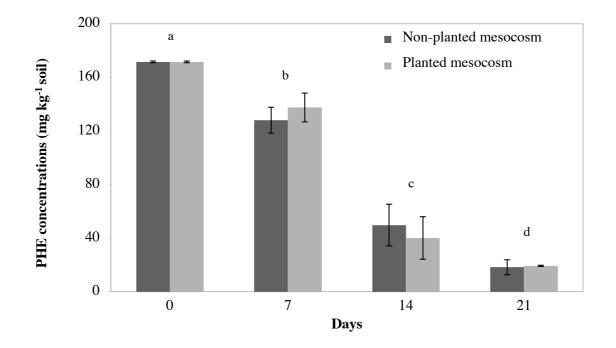


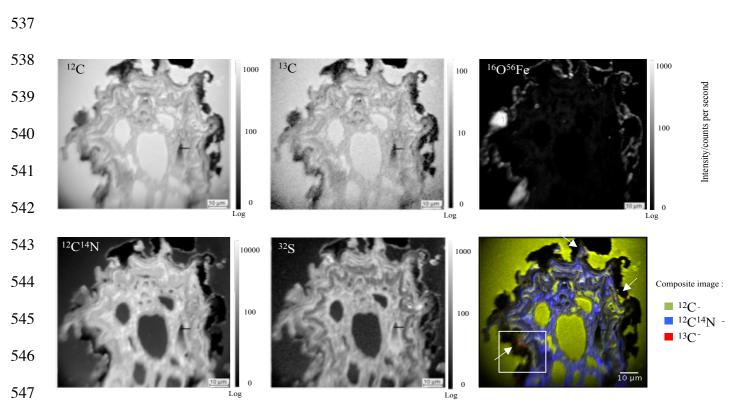


Figure 1. Extractable concentrations of phenanthrene (PHE, mg kg⁻¹ of soil) in planted and non-planted mesocosms at different time-points over a 21-day period. Means \pm SD, n = 4. Different letters indicate significant differences (p<0.05) between the sampling times.



526

527 Figure 2. ¹³C dynamics after ¹³C-PHE spiking in non-planted (left) and planted (right) 528 mesocosms at different time-points over a 21-day period. (a) ¹³C enrichment (E¹³C, at.%) in the 529 plant and soil compartments. Means \pm SD, n = 4. ¹³C enrichment was calculated from δ^{13} C data 530 obtained by EA-IRMS. (b) ¹³C distribution in the mesocosm compartments (PAH extract, 531 residual soil, plant, and C loss). Results are expressed as percentages of the ¹³C-PHE 532 concentrations initially added in the soil. The plant compartment gathers root and shoot data. 533 PAH extract concentrations were calculated as the difference between bulk and residual soil 534 concentrations.



548

Figure 3. NanoSIMS elemental maps of ¹²C⁻, ¹³C⁻, ¹²C¹⁴N⁻, ³²S⁻ and ⁵⁶Fe¹⁶O⁻ from labeled ryegrass roots after 14 days. The 90*90- μ m² images are cross views of whole roots with adherent soil particles. A composite image (bottom right) was obtained by overlaying the ¹²C⁻, ¹²C¹⁴N⁻ and ¹³C⁻ images. Three areas of interest (bordered in white) with higher ¹³C content were identified. NanoSIMS data were analyzed using Winimage software.

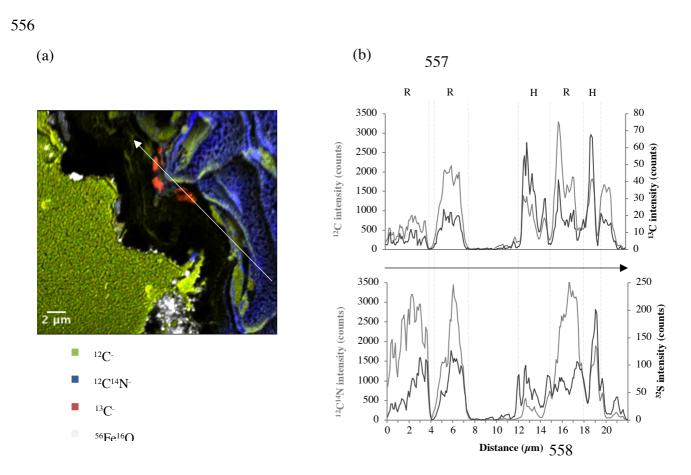




Figure 4. NanoSIMS composite image and linescan obtained from root tissue after 14 days. The zone is located inside the white square of Figure 3. (a) Composite $25*25-\mu$ m² image with hotspots reflecting the ¹³C-enrichment (red) obtained by overlaying ¹²C⁻, ¹³C⁻, ¹²C¹⁴N⁻ and ⁵⁶Fe¹⁶O⁻ (white) ion species images. (b) A white segment was drawn through these hotspots