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Abstract

The Jbel Haïmer copper-sulphide mineralisation occurs at 20 km north of Marrakesh in the Variscan Jebilet massif, Morocco. Most of the ores (up to 3.9 wt. % Cu, ≤ 38 ppm Ag, and up to 2.9 ppm Au) occur as impregnations of NE-SW fault/fracture zones and related tectonic breccia. Two independent stages of fluid circulation and mineral deposition are distinguished. First, Late Variscan high temperature-low pressure metamorphism synchronous of granite intrusion induced percolation of C-H-O-N hydrothermal fluids throughout faults and shear zones. The drop in pressure from lithostatic down to hydrostatic values at temperatures around 350 ± 50°C triggered quartz precipitation associated with minor Sn-As-(Co-Ni) deposits and finally brecciation. The second stage, more recent than Triassic, consists in the deposition of quartz + carbonates, followed by Cu-(Pb-Zn) sulphides (±Ag-Au alloys) in fractures crosscutting 240 Ma microdiorite dikes. Mixing of evaporitic brines likely coming from Triassic formations with low salinity aqueous fluids was responsible for Cu-sulphide deposition at temperatures of around 220-280°C at a depth of 4-5 km. These base metal-rich brines are similar to those from the nearby Roc Blanc Ag-deposit and several other silver and base metal deposits in Morocco, considered as having circulated during the Central Atlantic Ocean opening.

Keywords: Cu-(Ag) ores, brines, fluid mixing, Triassic evaporite, Atlantic rifting

Introduction

Granite-related ore systems, an essential target for the mineral exploration industry (Sial et al., 2011), were developed as models for several commodities such as Au and W-Sn (granite related Au-deposits (IRGDs) for instance: Thompson et al., 1999; Hart, 2007; Thompson and Newberry, 2000; Lang and Baker, 2001). The proximity of the deposits to intrusions or inferred intrusions marked by contact metamorphic aureoles was one of the main arguments for such ore deposit models where magmatism plays a significant role at less as a source of heat flow for fluid convection. In Morocco, base metal (Pb-Zn-Cu), and precious metal (such as Ag) deposits are however not coeval with Pan-African or Variscan orogenies and related granitoids. Thus, extensional tectonics associated with the successive stages of Atlantic rifting favoured the formation of Ag and Pb-Zn deposits in Morocco and Western Europe. Ore deposits formed either in the Mesozoic sedimentary cover or in the basement where inherited discontinuities are excellent pathways for sedimentary brines. There, the increased porosity and permeability of damaged zones developed around faults and shear zones favoured fluid flows (Boiron et al., 2010; Nadoll et al., 2019).

In the Variscan Jebilet massif (Morocco), several vein-type polymetallic deposits occur in close spatial association with Late Variscan plutons such as the Roc Blanc Ag-(Zn-Pb) deposit where Ag grades reach 500 – 700 ppm or the Jbel Haïmer Cu occurrences (Huvelin, 1977). Huvelin (1977) considered these deposits as related to magmatic intrusions. At Roc Blanc, the model recently proposed by Essarraj et al. (2017a) is very different because it considers that the introduction of silver is late and not related to the Variscan cycle but synchronous to the Atlantic rifting after Triassic times. The silver deposit was there superimposed on previous Variscan mineral assemblages (sulphides and sulpharsenides) and related to the downward circulation of basinal brines.

At Jbel Haïmer, located at most 500 metres from the Roc Blanc deposit, a series of copper occurrences have been mined before the XXth century only in the near-surface weathered part. Copper was exploited in metre-sized excavations as malachite in breccia affected by intense oxidation. Oxidised ore breccia from outcrops at Jbel Haïmer have Cu grades up to 3.9 % (average 0.6 %), up to 38 ppm Ag, and 0.5 ppm to locally 2.9 ppm Au (Krutilin et al., 2013). The average Cu grade of the Jbel Haïmer ore is similar to those from volcanogenic massive sulphide deposits that occur a few kilometres towards the South-West and North-West, such as the Draa Sfar deposit (0.5 wt. % Cu), and the Kettara abandoned mine (0.6 wt. % Cu). There, Cu-rich quartz-carbonate veins crosscut volcanogenic massive sulphide (Outigua et al., 2020; N'Diaye et al., 2016).

The present study aims to decipher the nature of the mineralising fluids in the Jbel Haïmer Cu-deposit, in particular the temporal and spatial relationships between the mineralisation and the geodynamic events. The objective was to test a potential generalisation of the Roc Blanc genetic model developed for Ag-ores to other base metals such as copper in the Central Jebilet.

Regional geology

Variscan units and tectonic-metamorphic stages: The Jebilet massif extends 170 km E-W and 7 to 40 km N-S (Fig. 1), and consists of three lithotectonic domains, the Western, Central, and Eastern Jebilet, separated by first-order shear zones: the Marrakesh Shear Zone (MSZ), between the Eastern and Central Jebilet (Lagarde and Choukroune, 1982) and the West Moroccan Shear Zone (WMSZ) between the Western and the Central Jebilet (Huvelin, 1977; Piqué et al., 1980; Lagarde and Choukroune, 1982; Le Corre and Bouloton, 1987) (Fig. 1b). According to Huvelin (1977) and Aarab and Beauchamp (1987), this massif is an intra-continental rift-basin filled with marine sediments deposited in an anoxic environment (Beauchamp, 1984; Beauchamp et al., 1991) from the Late Devonian to Early Carboniferous. During this extensional tectonic

93 regime, cogenetic tholeiitic mafic-ultramafic and calc-alkaline felsic intrusions are the
94 witnesses of the pre-orogenic bimodal magmatism (Bordonaro, 1983; Aarab, 1984) (Fig. 2a).
95 Recent U-Pb LA-ICP-MS dating on zircon from granodioritic plutons gave an absolute age of
96 358 ± 7 Ma (Tabouchent - Bamega) and 336 ± 4 Ma (Oulad Ouaslam) (Delchini et al., 2018).
97 This age range is older than the former ages obtained for the same rocks (330.5 ± 0.7 Ma, U-
98 Pb on zircons, Essaifi et al., 2003; and 327 ± 4 Ma, Rb-Sr on whole rock, Mrini et al., 1992).
99 These intrusions produced contact metamorphism (M1) in the pyroxene-hornfels facies
100 (Huvelin, 1977; El Hassani, 1980; Delchini et al., 2018). The Jbel Haïmer Cu-ore bodies occur
101 in the central part of the Jebilet comprised of the so-called Sarhlef series (Fig. 1). The Sarhlef
102 metasediments deposited in the rift-basin as a thick succession of argillites overlain by
103 carbonaceous argillites and limestones (Bordonaro, 1983). These series have a Carboniferous
104 age (Upper Visean - Namurian: fossil relative dating by *Posidonomya becheri* Bronn; Huvelin,
105 1961). In the Sarhlef series, stratabound lenticular Fe-Cu-Pb-Zn volcanogenic massive
106 sulphides formed during this rifting period (Huvelin, 1977). The structural features of the Jebilet
107 massif are due to the succession of two deformation events, D1 and D2 that lasted from the Late
108 Carboniferous to the Early Permian and constituted the Variscan orogeny (Delchini et al.,
109 2018). During the D1 compressive deformation, the emplacement of superficial allochthonous
110 nappes resulted in recumbent folds (Delchini et al., 2018), accompanied by rare schistosity
111 cleavage S1, boudinage of competent layers as well as vertical fractures and veins. The context
112 is typical of a diagenetic to anchizonal metamorphism transition (low temperature: 150-200°C,
113 Bamoumen, 1988). The D1 structures trending E-W in the Central Jebilet and the West of the
114 Eastern Jebilet suggest an N-S shortening (Delchini et al., 2018). The D2 is the major polyphase
115 deformation event that corresponds to a WNW-ESE to NW-SE transpressional crustal
116 shortening. During the first D2 stage, NS–N30°E kilometre-size isoclinal folds developed, as
117 well as a penetrative S2 cleavage and a high temperature–low pressure (HT-LP) thermal

metamorphism (greenschist to amphibolite facies; 4 to 5 kbar and 560-485°C; Delchini et al., 2016). The second stage of D2 deformation corresponds to transpressional shearing that induced shear zones at all scales, in particular the multi-kilometre-sized WMSZ and the MSZ (Fig. 1b). HT-LP thermal metamorphism M2, so-called M2a in Delchini et al. (2018), isoclinal folding D2, high thermal gradients (50°C/km, and up to 95°C/km) and shear zone activity occurred within the 310-280 Ma period (Delchini et al., 2018). Close to MSZ, leucogranites such as the Bramram intrusion (Fig. 2a) at 295 ± 15 Ma (Rb/Sr, Mrini et al., 1992) and 297 ± 6 Ma (Tisserant, 1977) intruded the Sarhlef schists. Local contact metamorphism (noted M2b in Delchini et al., 2018) reached hornblende-hornfel facies with peak conditions about 2–3 kbar and 620°C (El Hassani, 1980; Delchini et al., 2016) and locally induced the development of graphite-rich skarns (Fig. 2a) in the Central Jebilet (Bastoul, 1992). The HT-LP metamorphism coeval with leucogranite intrusions may be explained by the inversion of the thermally weakened intracontinental domain during the Variscan compression D2 (Delchini et al., 2018). A final increment of the D2 stage corresponds to the development of a network of regional brittle conjugate shear zones and brittle reverse faults (ENE, and NW) that locally deform the S2 major cleavage (Delchini et al., 2018). The Aït Bella OuSaïd kilometre-sized fault located close to Jbel Haïmer (few hundred metres to the North, Fig. 2a) is one of the longest and thickest ENE shear zones resulting from the last D2 stage (Delchini et al., 2018).

Post-Variscan microdiorite dikes and sills: they crosscut both the bimodal magmatic and granodioritic rocks and leucogranite plutons. They may exceed 1 km in length, strike N-S to N40°E (Fig. 2a). They formed during a period of distension (Huvelin, 1977; Bordonaro, 1983; Aarab and Beauchamp, 1987; Bernard et al., 1988; Bouloton and Gasquet, 1995; Essaifi et al., 2003; Dostal et al., 2005) at ca. 240 ± 10 Ma (K-Ar method on kaersutite: Gasquet and Bouloton, 1995; Youbi et al., 2001; Dostal et al., 2005, summarised by Bouloton et al., 2019). They are synchronous with the thermal events which affected the Western Meseta in the 250-

143 210 Ma range (apatite fission track thermal modelling: Ghorbal et al., 2008; Saddiqi et al., 2009;
144 Barbero et al., 2011).

145 *Mesozoic sedimentation and deformation:* several extensional deformation events occurred
146 from lower Triassic to Post-Kimmeridgian times. Such events are coeval with the break-up of
147 Pangea and Central Atlantic opening (Medina, 1991; Jalil, 1999; Tourani et al., 2000; Medina
148 et al., 2001; El Arabi, 2007). Subsequently, the deposit of thick clastic series alternating with
149 evaporites and basaltic flows, filled large sedimentary basins. In this time, the Central Jebilet
150 domain was a part of the West Moroccan Arch (WMA: Hafid, 2006; El Arabi, 2007; Saddiqi
151 et al., 2009). The latter was a paleogeographic high separating Tethyan domain to the East, i.e.,
152 the High Atlas rift and the Atlantic rift to the West. Paleogeographic reconstructions and apatite
153 fission-track data (Saddiqi et al., 2009) showed that the Central Jebilet high was supplied by
154 sediments that did not exceed 2.4 km in thickness, an estimate slightly lower than that proposed
155 by Ghorbal et al. (2008). Triassic series, composed of sandstone, argillites, and evaporites
156 (gypsum and salt), are outcropping at the Southern limit between Central and Eastern Jebilet
157 (Fig. 2a) as well as in the Western Jebilet (Huvelin, 1977). The pre-Atlasic orogenic period
158 lasted from the Triassic to the Late Cretaceous. From a structural point of view, the parallelism
159 of some major Paleozoic trends (NNE, High and Middle Atlas) attests to the reactivation of
160 former structures during the Atlasic orogeny at the Eocene (Frizon de Lamotte et al., 2008).
161 Nevertheless, the E-W striking fault systems that crosscut older Paleozoic structures in the
162 Jebilet did not correspond to reactivated structures (Frizon de Lamotte et al., 2008). Those
163 newly formed E-W fracture sets are filled with quartz-carbonate and contain Cu and Pb-Zn-
164 (Cu) ores (Huvelin, 1977).

165 *The Jbel Haïmer Cu-ore deposit*: At Jbel Haïmer, Cu ore bodies (Fig. 2b) are hosted by
166 greenish-brown to grey Sarhlef schists, split in millimetre thick sheets (S1). Intercalations of
167 sandy schist and sandstone layers prevail as NE-SW protruding crests at most one metre thick

and a few tens of metres long. Grey quartz occurs as boudinaged lenses in the schistosity planes, derived from metamorphic exsudation. Sarhlef series sometimes present hornfels with cordierite spots produced by contact metamorphism during granodiorite intrusions (Huvelin, 1977). Granodiorites outcrop a few kilometres to the South of Jbel Haïmer (Fig. 2a). Besides, contact metamorphism locally transformed sandstones to quartzite. The latter shows granoblastic mosaic texture and presents millimetre-sized impregnations and veinlets of dark-brown tourmaline and hyaline quartz that crosscut the schist. When approaching granodiorite, tourmaline abundance increases up to massive black tourmalinite. Several microdiorite dikes (< 1m in width and up to one hundred metres in length) are widespread in the Jbel Haïmer outcrops and throughout the mineralised areas (Fig. 2a).

Ore zones occur exclusively in breccia affecting competent formations such as sandy schist and sandstone - quartzite layers (Fig. 3a to 3c). The old mining works are no longer accessible in Jbel Haïmer and were confined to a few tens of metres depth over the water table (up to 50 m, Huvelin, 1977). Most ores are therefore malachite and Fe-oxide impregnations in the oxidised breccia (Fig. 3a to 3e) and locally in the brecciated microdiorite dikes (Fig. 3f, 3g).

Methods

Representative samples were collected from the wall rocks (spotted schist, sandy schist-sandstone, and microdiorite dikes), as well as from the ore breccia and quartz veins for this study. Petrographic and scanning electron microscope (SEM) investigations were carried out using twenty selected samples.

Nine doubly polished wafers were used for fluid inclusion studies. Fluid inclusion typology (Table 2) follows notations from Boiron et al. (1992): aqueous (w), carbonic (c), bulk homogenisation to the liquid phase (L), bulk homogenisation to the vapour (V), presence of halite cube (h), and unidentified solids (s). Microthermometry was performed using a Linkam

193 THSMG600 freezing-heating stage (Shepherd, 1981). Abbreviations follow the convention
194 from Roedder (1984) for microthermometric data presentation: Tm CO_2: melting temperature
195 of volatile phase for carbonic fluid inclusions, Th CO_2: homogenisation temperature of volatile
196 phase for carbonic fluid inclusions, Te: eutectic or apparent eutectic temperature defined as the
197 temperature at which the first liquid appears in the frozen fluid inclusion during low-
198 temperature measurements, Tm hh: temperature of final melting of hydrohalite, Tm ice: final
199 melting temperature of ice, Tm cl: melting temperature of clathrate, Th (L or V): total
200 homogenisation temperature (L+ V to liquid or L + V to vapour phase respectively) and Tm
201 NaCl: melting temperature of halite cube). The precision of phase transitions below 0°C was
202 +0.2°C, whereas the precision of the liquid-vapour homogenisation was 1°C. Over the range –
203 90 to 0°C, the accuracy was 0.4°C, and from 0 to 300°C was 2°C. For most aqueous fluid
204 inclusions, salinity calculations have been made in the H_2O-NaCl system using data from
205 Bodnar and Vityk (1994) based on Tm ice or Tm halite for fluid inclusions containing halite
206 cube (when Tmh < Th and no Tm hh is available). When both Tm ice and Tm hh or Tm NaCl
207 values were available, salinity was calculated in the H_2O-NaCl-$CaCl_2$ system using data from
208 Steele-MacInnis et al. (2011).
209 Molar fractions of gas components (CO_2, CH_4, N_2) were determined on individual carbonic and
210 aqueous-carbonic fluid inclusions in quartz by Raman spectrometry analysis using a DILOR
211 LABRAM Raman spectrometer at the GeoRessources laboratory, University of Lorraine,
212 Nancy, France. Bulk composition, molar volume, volatile/water ratio, and salinity were
213 calculated from the *P-V-T-X* properties of individual fluid inclusions in the C-O-H-S system
214 (Dubessy, 1984; Dubessy et al., 1989; Thiery et al., 1994; Bakker, 1997). Isochores were
215 calculated using the program ISOC from the computer package FLUIDS-1 (Bakker, 2003).
216 Laser ablation - inductively coupled plasma - mass spectrometry (LA-ICP-MS) analyses have
217 been performed at the GeoRessources Laboratory (Nancy, France) on individual fluid

inclusions to quantify the ionic composition of fluids involved in the metal deposition. Laser ablation was carried out with an ESI New Wave Research UC 193 nm excimer laser at 5 Hz laser frequency. The ablated material was analysed with an Agilent 8900 Triple Quadrupole ICP-MS, equipped with a collision–reaction cell. The ablated material was carried in helium gas (0.5 l min-1), which was mixed with argon (0.9 l min-1) via a cyclone mixer (volume of 9.5 cm^3) before entering the ICP torch. A straight ablation technique with spot sizes of 30 and 40 μm was applied to obtain overall higher signal-to-background intensity ratios and lower limits of detection (LOD). Each analytical series consists in the following isotopes: ^{23}Na, ^{24}Mg, ^{39}K, ^{44}Ca, ^{55}Mn, ^{57}Fe, ^{63}Cu, ^{66}Zn, ^{75}As, ^{85}Rb, ^{87}Sr, ^{109}Ag, ^{121}Sb, ^{137}Ba and ^{208}Pb. The collected intensity ratios for Na were converted to concentration ratios by external calibration against a NIST 610 standard reference glass following a bracketing standardisation procedure (Longerich et al., 1996). ^{23}Na was used as an internal standard to calculate absolute concentrations and was obtained via microthermometry. The calculation method for the chemical composition of the fluid inclusions is based on a combination of the ice-melting temperature, a Pitzer-based thermodynamic model, and the LA-ICP-MS Na/cation ratios (Leisen et al., 2012a). Limits of detection for a given element can be calculated, according to Longerich et al. (1996) and described in Leisen et al. (2012a).

Barren and ore veins in their geological context

The schistosity cleavage (S1) in the Sarhlef series strikes N30-50°E with an average dip of 20-60° to the East (Fig. 4a, 5a). Joints and faults are mostly parallel to the schistosity or crosscutting it (Boutira and Hakkou, 2012; Essarraj et al., 2013; Fig. 4, Fig. 5).

Barren veins: fractures, a few centimetres to a few metres in length, form a dense network striking NW-SE (average N130-140°E), dipping 50° (East and West) to sub-vertical. Pale grey quartz QI infillings are found in fractures crosscutting competent sandy schist and sandstone

and are associated with chlorite along the vein walls (Fig. 4b, 5b to 5e). The second set of fractures and white quartz veins named QII, in general up to a few metres long, strikes from N160°E to NE-SW and crosscuts the NW-SE QI veins (Fig. 5d) and fractures. The veins show a geodic texture locally. Decametre length corridors of breccia are oriented N10-30°E to NE-SW (Fig. 3a to 3e, 4) with an average dip around 60° to the East, and affect all preceding quartz types. They correspond to open structures sub-parallel to the strike of the regional schistosity, although schistosity planes are difficult to be observed in sandy schist and sandstone layers (Fig. 3a, 3c). Microdiorite dikes intruded the schist parallely to the ore breccia corridors, as they strike mostly NNE-SSW (Fig. 4). They are locally deformed in particular along the breccia corridor (Fig. 3g).

Quartz-carbonates veins and Cu-ore bodies: a set of E-W fractures, feathery quartz and carbonate veins crosscut the schist and sandstone layers, as well as the microdiorite and most earlier veins above described (Fig. 5f, 5g). The Cu- ores formed later than E-W fractures and crosscut as well as the microdiorite. They are found as infillings of new microfracture sets in the vicinity of the main ore breccia corridor.

Mineralogy

Table 1 provides the paragenetic succession.

Pre-ore stage: A biotite-tourmaline-pale-grey quartz QI assemblage constitutes the first stage of barren vein infillings. Biotite, with tourmaline crystals, are abundant and consistently line the QI vein walls (Fig. 6a, 6b). Biotite was almost totally replaced later on by chlorite (chlorite 1). The pale-grey quartz QI is coarse-grained and intensively sheared. It shows several interlocked networks of microfractures (fluid inclusion planes, FIPs) mainly NW-SE and NE-SW. Cassiterite, as crystals up to 300 μm in size, often with Sb traces, and sulphides (pyrite, arsenopyrite, and Co-As-Ni minerals, bismuthinite with Se traces) are found in brecciated QI,

as well as small disseminated grains (a few μm) of native Bi (traces of Se and Ag) (Fig. 7a). The coarse-grained white quartz QII is locally associated with brecciated tourmaline agglomerations and followed by pyrite (Fig. 6c).

Fe-Co-Ni sulpharsenides fill fractures or cavities of the breccia and postdate pyrite and QII. They consist of euhedral arsenopyrite, Fe-(Co)-sulphide (Co-pyrite with around 6 wt % Co), and small size (few μm) euhedral crystals of Fe-Co-Ni-sulpharsenides and cobaltite (Fig. 7b).

After the breccia stage, chlorite 2 and QIII formed around the breccia fragments. Locally, pyrrhotite is spatially associated with chlorite 2 and quartz QIII. In schist, millimetre size clear quartz (QIII) veins are observed sub- parallely to the QI veins, and are lined by a newly crystallised chlorite 2 (Fig. 6d, 6e). They correspond to the re-opening of the QI vein selvage (Fig. 6e). In these zones, QI-QII quartz crystals present recrystallisation halos or epitaxial clearer growth bands at the contact of chlorite 2.

Cu-ore stage: quartz breccia and microdiorite are crosscut by millimetre to centimetre-sized fractures, lined by feathery comb quartz, siderite and calcite (Fig. 6f to 6h). The late quartz (QIV) is clear, undeformed, and presents euhedral geodic terminations. It hosts scarce fluid inclusions and fluid inclusion planes (Fig. 6i). Chalcopyrite occurs later than QIV as impregnations of the former breccia and microfracture infillings. It sometimes contains traces of Te, scarce inclusions of Fe-sphalerite (up to 8 wt % Fe), stannite as well as an unidentified Bi-Te-Se-S mineral (Fig. 7c, 7d and 7e). Galena is found sporadically in microfractures affecting earlier Fe-As minerals (Fig. 7b and 7f). Selenium rich-galena (up to 3 wt % Se; traces of Bi) locally crosscuts chalcopyrite (Fig. 7g). Precious metals are present as late Ag-Au alloys (Ag/Au around 2).

Supergene alteration: supergene minerals replace sulphides and impregnate breccia fragments and microfractures (Fig. 6c, 6d, and 7h). Thus, exploited copper ores are mostly malachite (Fig. 7c, 7e, and 7g) that results from the weathering of chalcopyrite from the Cu-ore stage. Fe–

oxides/hydroxides are predominant and are accompanied by a few grains of covellite and cerussite. Native metals (Pb, Zn, Ag, Cu, Bi, Se) are present as minute grains (10 to 30 μm).

Fluid inclusion study

Fluid inclusions (FIs) were studied in quartz and carbonate from the mineralised breccia as well as in most fracture quartz infillings in the Jbel Haïmer area. Fluid inclusions types are defined in Table 2. Table 3 provides microthermometric data and Table 4 Raman data.

Pre-ore stage

Fluids in tourmaline: primary FIs are two-phase and water dominated with traces of a low-density volatile, noted Lw-(c)-tur (Table 3). Their salinity is estimated around 24 – 26 wt% NaCl equiv. Th, to the liquid phase, ranges between 370 and 390°C. Secondary FIs in quartz from tourmalinised quartzite are two-phase water dominated FIs with traces of low-density volatile (Lw-(c)1), salinity ranging between 15.7 to 16.5 wt% NaCl equiv. and Th from 286°C to above 386°C to the liquid phase.

Fluids in early quartz Q1 – biotite/chlorite 1 stage: quartz QI in NW-SE veins and quartz fragments in breccia contain dense clouds of FIs (Fig. 8a) where true primary FIs cannot be easily distinguished. Two types of FIs distribute randomly in QI crystals: i) water dominated FIs with traces of volatiles (Lw1), which are abundant in FIs clouds, have a salinity ranging from 20 to 26 wt% NaCl+$CaCl_2$ equiv. and Th from 296° to over 357°C to the liquid, some of them decrepitating at temperatures above 350°C; ii) one-phase dense carbonic liquid inclusions, Lc1, found dispersed in QI show Tm CO_2 from –78.4 to –67.9°C, indicative of the presence of significant amounts of other volatiles. Th CO_2 to the liquid phase ranges from –70.2 to –13.7°C with a mode at –32°C (Table 3).

Lw2 aqueous FIs containing relatively high vapour infilling (up to 30%), with small amounts of gas (low density of the volatile phase), are abundant as regular networks of FIPs crosscutting

QI quartz crystals from veins and breccia. The average salinity ranges from 11.8 to 15.8 wt % NaCl equiv. and Th ranges from 261 to 317°C (Table 3).

Fluids in quartz QII (NS to NE-SW white quartz veins): scattered one-phase carbonic vapours (Vc2) and two-phase aqueous carbonic FIs (Lc-w2) coexist in the same QII microdomains (Fig. 8b, 8c). In both FI types, Tm CO_2 ranges from –102.8 to –68.5°C and Th CO_2 from –71.6 to –60.5°C to the vapour phase (Table 3). The volatile phase is dominated by N_2 (60-70 mol%), (Table 4). Tm cl and Tm ice for Lc-w2 are respectively from –4.2 to 1°C and from –8.4 to –3.8°C. Th of Lc-w2 FIs are recorded over 460°C to the liquid phase, but most of the FIs decrepitate around 350°C. In QII veins, one-phase carbonic vapours (Vc2) and two-phase aqueous-carbonic inclusions (Lc-w2) coexist in the clear quartz microdomains surrounding pyrite. Vc2 show Tm CO_2 around –102°C and Th CO_2 around –97°C while Lc-w2 recorded Tm ice around –10 to –9°C, Tm cl from 6.2 to 9.1°C and Th from 414 to 453°C to the liquid. Raman spectroscopy analyses show that the volatile phase of Lc-w2 FIs in recrystallised QII is dominated by CH_4 (60-70% mol%), (Table 4).

Vapour FIs showing roughly similar microthermometric characteristics to Vc2 (Tm CO_2 modes at –85°C and –65°C and Th CO_2 at –65 to –60°C to the vapour) are present as FIPs crosscutting QI. They are random FIs in recrystallised QI.

Fluids in quartz from breccia and NW-SE QIII – chlorite 2 veinlets: carbonic liquids (Lc3) (Fig. 8d to 8h) are monophase FIs scattered and pseudo-secondary in clear quartz QIII surrounding chlorite 2 crystals and in QIII from veins (Fig. 9). Lc3 form also sets of FIPs parallel to vein walls in QI (Fig. 8d and 8e) and crosscutting QII. Tm CO_2 ranges from –65 to –57.6°C, and Th CO_2 goes from –46.9 to 18.6°C to the liquid phase (Table 3). The Lc3 FIs pseudo-secondary in QIII (Fig. 8h) display a Tm CO_2 close to those of pure CO_2 (Tm of pure CO_2: –56.6°C) and Th CO_2 close or above 0°C. CO_2 dominates Lc3 FIs (59 to 88 mol%), followed by N_2 (10 to 29 mol%) with lower CH_4 content (2 to 14 mol%) (Table 4, Fig. 9).

343

344 ***Cu-ore stage fluids***

345 Feathery quartz, calcite, and QIV are devoid of carbonic FIs and show exclusively aqueous FIs
346 covering a broad range of salinities. The two major FI types are (Fig. 10, Table 3): i) high
347 salinity FIs with a halite crystal and an unidentified clear solid (noted Lwh), and sometimes a
348 few additional solids (Lws) and ii) moderate to low salinity FIs (Lw) (Fig. 10, Table 3).
349 Chalcopyrite is present in the same quartz microdomains as Lwh brines as well as within Lw
350 FI assemblages from clear quartz QIV microdomains.

351 Almost all of the aqueous FIs present relatively low eutectic temperatures (–65°C), likely
352 corresponding to metastable eutectic of the H_2O-NaCl-$CaCl_2$ system (Davis et al., 1990). These
353 FIs are present as FIPs in early quartz (QI, QII, and QIII).

354 *High salinity brines*: Lws FIs have an irregular shape (Fig. 10a to 10c), and distribute in
355 recrystallised microdomains of quartz QI-QII from veins and breccia, as well as irregular FIPs
356 crosscutting QI, QII, and QIII (Fig. 8g). Salinity is around 38 - 39 wt% NaCl+$CaCl_2$ equiv.
357 Most Lws FIs decrepitate upon heating at temperature from 230°C with an average
358 decrepitation temperature of 300-330°C. A few Th were recorded between 174 and 252°C to
359 the liquid phase. The dissolution of solids occurs from 90 to 360°C, mostly at a temperature
360 higher than Th.

361 Lwh FIs present regular shape and sometimes one unidentified rounded or rhombic-shape clear
362 solid in addition to the halite cube (Fig. 10d and 10g), and occasionally a small opaque solid.
363 Lwh FIs are scattered in QIV and form abundant FIPs crosscutting QI, QII, and QIII veins and
364 previous volatile-rich FIPs (Fig. 10e). Besides, Lws and Lwh FIs are sometimes present in the
365 same recrystallised early quartz microdomains where Lwh FIPs often crosscut Lws FIPs.
366 Salinity ranges from 31.3 to 37.3 wt% NaCl+$CaCl_2$ equiv. (average 34 wt% NaCl+$CaCl_2$
367 equiv.). Th ranges from 179 to 264°C (mode around 220°C), and halite cube melts at

368 temperatures ranging from 155 to 262°C, the second solid remaining unmelted. Some FIs have
369 decrepitated around 270°C. For most Lwh FIs, Tm NaCl and Th occur roughly in the same
370 average domain between 220 and 230°C.
371 Lw FIs in feathery quartz and calcite show Tm ice from –32 (primary in calcite) to –20°C
372 (secondary in feathery quartz and calcite) and Th from 186 to 242°C. Salinity ranges from 22.4
373 wt% to > 30 wt% NaCl equiv.

374 *High to low salinity fluids*: several types of two-phase aqueous FIs (Lw), sometimes with an
375 unidentified clear solid and sporadically an opaque solid, form regular networks of FIPs in QI,
376 QII, QIII and QIV (Fig. 10g and 10h). Most aqueous FIs show a regular shape and have a
377 vapour infilling around 10%. Tm ice range from –29.6 to –7.1°C, Tm hh range between –11.9
378 and –6.5°C and Th range from 139 to 249°C with several modes corresponding to different
379 FIPs. Salinities range from 10.6 to 28 wt% NaCl+$CaCl_2$ equiv. Clear solids do not melt until
380 temperatures over 330°C. FIs in planes crosscutting QIV crystals show salinities from 9 to 15.3
381 wt% NaCl equiv. and Th from 156 to 242°C. Scarce FIs, scattered in the recrystallised quartz
382 around chalcopyrite (Fig. 10f), have a salinity around 23 wt% NaCl equiv. and Th around 150-
383 160°C.
384 A general trend corresponds to a decrease in salinity at constant or slightly decreasing Th from
385 the Lws brine to Lw FIs (Fig. 11). The salinity evolution is in agreement with chronological
386 relationships between aqueous fluid inclusion types that start with Lws, crosscut by Lwh FIPs,
387 and both of them being crosscut by several Lw FIPs that show a gradual decrease in salinity.
388 The latest Lw FIs that crosscut quartz QIV present the lowest salinities (Fig. 11). Chalcopyrite
389 deposition occurred when mixing occurred, e.g. when fluids were characterised by intermediate
390 features between the two end-members of this general trend (i.e., Lws and Lw FIPs in QIV).

391 *Major elements and metal content of the brine inclusions*: LA-ICP-MS data for Lws and Lwh
392 brines are given in Table 5 and Fig. 12. Lw fluid inclusions were too small to be analysed.

Lws brines (FIPs in QI, QII and QIII) show relatively high Na content (average 2735 mmol/kg), Ca, and K, (average around 941 mmol/kg,and 397 mol/kg respectively) but low Mg content (73 mmol/kg) (Table 5, Table 6). Average molar Na/Ca, Na/K, and Na/Mg ratios are 2.9, 6.9, and 37.7, respectively. Lws FIs display Pb and Zn contents of 72.2 mmol/kg and 159.1 mmol/kg respectively on average and a Cu content of 19.1 mmol/kg. Ag concentration is around 0.2 mmol/kg (Table 5, Table 6).

Lwh FIs (quartz QIV): in Lwh FIs, Na is also the dominant cation (average around 1532 mmol/kg), but they are rich in Ca (1174 mmol/kg), K and Mg (580 mmol/kg and 501 mmol/kg, respectively) (Table 5, Table 6). The average molar ratios are around 1.3 for Na/Ca, 2.6 for Na/K and 3.1 for Na/Mg. High amounts of Pb, Zn and especially Cu are present in Lwh brines (average: 28.8 mmol/kg, 136.8 mmol/kg and 43.9 mmol/kg respectively) as well as Ag (average. 5.3 mmol/kg) (Table5, Table 6). In Lwh brines, Cu and Pb concentrations are correlated with Na. Fe is detected at high concentrations, and is certainly present. Fe contributions were however not considered in the calculations because of probable contaminations by solids yielding unrealistic contents.

From Lws to Lwh brines, K, Ca, Mg, Cu, As and Ag concentrations increase, together with a decrease in Na, Zn, Pb, Ba, Rb, and Sr concentrations (Fig. 12, 13). Similar positive correlations are obtained for Sr–Ba, and As–Cu. By comparison, brines from Jbel Haïmer contain higher Zn, Cu, Pb and Ag contents than those from the Roc Blanc deposit (Essarraj et al., 2017a), (Table 6 and Fig. 13). The high chlorinity, the richness in K, Ca and Mg are typical of evaporitic brines as discussed below (Fig. 14).

Discussion

Pre-ore stage fluids in the regional metamorphic context

Early tourmaline-biotite-Sn assemblage: Tourmaline formed first from hot brines (around 25

418 wt% NaCl equiv., T > 380°C) related to granitoid intrusions, which produced hornfels in the
419 contact metamorphic aureole where tourmaline-rich pegmatite veins and lenses are abundant.

420 C-H-O-N fluids: NW-SE-striking fractures with quartz QI (biotite) infillings followed by NE-
421 SW fractures with quartz QII, formed from N_2-CH_4-(CO_2-H_2O) fluids at a temperature ranging
422 from 450 to 500°C. Brecciation occurred in the competent NE-SW sandy schist and sandstone-
423 quartzite layers later than QI, QII veins and is synchronous with new NW-SE fractures likely
424 conjugated with NE-SW fractures. Large circulation of CO_2-(H_2O) dominated fluid produced
425 the cementation of the breccia and newly formed NW-SE fractures filled with quartz QIII and
426 chlorite 2.
427 An evolution in the volatile phase composition (Fig. 9) occurred from N_2-CH_4 (± H_2O) rich
428 fluid associated with early quartz (QI, QII) towards CH_4-CO_2 (± H_2O) rich fluid associated with
429 pyrite deposition. Finally, CO_2-rich fluid associated with chlorite 2 deposition ends the volatile-
430 rich fluid series. Aqueous-carbonic fluids (N_2-CH_4-(CO_2 ± H_2O)) from Jbel Haïmer are very
431 similar to those observed in the Roc Blanc and Koudia El Hamra silver-(base metal) deposits
432 (Essarraj et al., 2017a; Nshimiyimana et al., 2018, respectively) (Fig. 9). Bastoul (1992)
433 documented fluids with similar volatile composition in skarns located a few kilometres North
434 of Jbel Haïmer (Fig. 9). Most volatile-rich fluids were produced during the HT-LP M2a stage
435 under regional metamorphism conditions (500-650°C and 200 MPa ± 20 MPa) (Bastoul, 1992;
436 Delchini et al., 2018). The production of H_2O-CO_2-CH_4 rich fluids in this P-T range results
437 typically from water and graphite equilibrium under reduced conditions close to Q-F-M buffer
438 which constitutes a rather good proxy of the redox conditions in metamorphic series (Huizenga,
439 2001; Huff and Nabelek, 2007). Additional local sources are needed to explain the mixing
440 trends among volatiles: i) a mixing trend between N_2 rich fluids and CH_4 rich fluids, and ii) a
441 mixing trend between a CO_2 rich fluid end-member and the N_2-CH_4 mixed fluids. Bastoul
442 (1992) found the same trends and attributed the formation of N_2 to the oxidation of NH_4^+ from

micas as also documented by Wright et al. (2012). Both authors have reported the occurrence of ammonium-rich phyllosilicates in metamorphic series and the related production of N_2 during their heating. N_2-CH_4 rich fluids are thus produced in graphite-rich series submitted to contact or retrograde metamorphism with highly variable N_2/CH_4 ratio and are in general very poor in water.

High thermal gradients: High temperatures recorded by early metamorphic fluids at Jbel Haïmer (Th > 450°C, this study) are in agreement with several temperature estimates: i) those proposed for the Central Jebilet (Bastoul, 1992), and ii) the temperatures derived from the graphite Raman thermometry (based on the structural organisation of the carbonaceous material to graphite) applied to hornfels (475 to 630°C, Delchini et al., 2016). Isotope data (O, H) of N_2-CH_4-CO_2 ± H_2O fluids either at the Roc Blanc deposit ($\delta^{18}O$ of waters in equilibrium with early quartz: 12.6 - 14.0‰ $_{V\text{-}SMOW}$; δD: –43.2 to –37.1‰ $_{V\text{-}SMOW}$, Essarraj et al., 2017a) or the Kettara deposit ($\delta^{18}O$ of waters in equilibrium with chlorite: 6.0 to 7.2‰ $_{V\text{-}SMOW}$; δD: –14.5 to –10.5‰, N'Diaye et al., 2016), compared to the composition of metamorphic waters from Sheppard (1986) confirm their metamorphic origin. Whatever their pristine source, a significant part of the described waters deeply equilibrated with metamorphic series, in particular with graphite producing H_2O-CO_2-CH_4 fluids. Early metal extraction and transport (Fe, Sn, As, Bi, Co, Ni) at Jbel Haïmer were probably favoured at that stage by aqueous-carbonic fluids set in motion in close connection with the HT-LP thermal metamorphism (M2a) and the late granites intruding the Sarhlef series (M2b). Aqueous-carbonic fluids play a significant role in metal transport and deposit in a great variety of Variscan deposits (Wilkinson et al., 1990; Cathelineau et al., 2020). N_2-rich fluids may be the main ore fluids in scheelite deposits (Wang et al., 2018) and CH_4-rich fluids in C-rich formations close to Au-As deposits (Guedes et al., 2002; Boiron et al., 2003).

P-T evolution: the pressure and temperature of trapping were estimated using isochores for most fluid types (Fig. 15A). The earliest fluids, Lw-(c)-tur FIs in tourmaline, were trapped at conditions close or slightly below those estimated for the thermal metamorphism (M2a). A significant temperature increase of approximately 250-300°C was therefore reached from regional metamorphic conditions, leading to thermal gradients of around 70°C/km (P-T box from Bastoul, 1992, Fig. 15A). Their isochores crosscut the P-T box for the HT-LP metamorphic conditions, which constitutes the maximal estimate. Isochores for the earliest FIs in quartz QI (Lw1) are subparallel to those of tourmaline FIs, while Lc1 isochores define a broad pressure range in between 70 and 150 MPa. Both fluids are considered as sub-synchronous. The intersection of their isochores provides an estimated pressure range of 70 to 150 MPa, and temperatures ranging from 360 to 460°C. For the second stage of quartz QII, consideration of the assemblage Lc-w2 and Vc2 yield to the following P-T estimate: 70-110 MPa, 450-520°C. The values are close to those of unmixing defined by the isopleths in the H_2O-CO_2-NaCl system for two different CO_2 contents. Such a process can explain the presence of Vc2 fluids. The third stage QIII is characterised by carbonic liquids that do not provide strong constraints on their trapping P-T box except the temperature estimate for the chlorite (300-350°C), which are similar to those of the Roc Blanc deposit. Carbonic liquids are thus trapped under higher pressure than in the preceding stage, in between 120 and 170 MPa.

The evolution of fluids from the tourmaline stage to QI and QII suggests a significant pressure decrease from lithostatic conditions towards hydrostatic conditions, in agreement with the pressure release expected after the breccia formation. During stage QIII, pressure increased again up to values close to those of the lithostatic regime. In both cases, the P-T box is intersected by a thermal gradient of around 60 and 70°C/km indicative of high thermal gradients in agreement with the thermal metamorphism and late intrusions of magmatic bodies defined by Delchini et al. (2018).

Most fluids from the early stages are therefore related to both abnormal thermal fluxes linked to HT-LP metamorphism and late granite magmatism during deformation stage D2. The presence of drainage zone, such as the MSZ and the Aït Bella OuSaïd shear zone, favoured collect and mixing of the different types of volatile-rich fluids generated during the thermal event (Fig. 16A). Fluids under lithostatic pressure may have produced brecciation and subsequent fluid decompression and precipitation of quartz QIII.

Cu brine circulation during post Variscan events in the Jebilet massif

P-T conditions: the intersection between the isochores calculated for the three FI types (Lws, Lwh, and Lw), and the thermal gradients expected in basins during extensional stages linked to rifting (see discussion below), provides an estimate P-T pair of 40-50 MPa and 220-280°C (Fig. 15B). Considering hydrostatic conditions and abnormal thermal gradient of 40°C/km and 50°C/km during this particular period, pressures correspond to a depth of around 4 to 5 km. Such a depth comprises the thickness of the Mesozoic series in the Jebilet (up to 2.4 km), and therefore yield to an estimate of the depth below the unconformity of 1.6 to 2.6 km.

Triassic formations as brine source: The Jbel Haïmer brines are characterised by high salinities from 31 to 37 wt% NaCl+$CaCl_2$ equiv. (average 34 wt% NaCl+$CaCl_2$ equiv.), are trapped around 220-280°C, and are enriched in K, Mg and Ca. They are similar to those of the Roc Blanc Ag-deposit (20 wt% to 35 wt% NaCl equiv., trapping temperature around 200°C and pressure of 50 MPa, Essarraj et al., 2017a) and the Koudia El Hamra Ag-deposit (Nshimiyimana et al., 2018).

The major element contents of the two brines Lwh and Lws were compared to those of seawater and evaporated seawater after the halite saturation. The high chlorinity, the richness in K (and Mg) are typical of brines having passed halite saturation (Fontes and Matray, 1993) (Fig. 14). They are similar to those described for evaporitic brines found either in a variety of sedimentary

basins or in basement faults (Poitou high (France): Boiron et al., 2002; Black Forest (Germany): Walter et al., 2019; Catalan coastal range (NE Spain): Piqué et al., 2018). The calcium enrichment, common to fluids from all sites, is generally interpreted as the result of water-rock interactions such as albitisation (Black Forest, Walter et al., 2016), and the relatively broad range of Mg contents as the result of Mg uptake by the crystallisation of dolomite in the sedimentary cover, and talc or other Mg silicates in the basement (Trimouns (Pyrénées, France): Leisen et al., 2012b).

Evaporites within the sedimentary clastic sequences, up to 8 km thick, from Atlasic basins, constitute good candidates for the source of brines. These subsident basins formed in relation with the Early Triassic to Early Jurassic extension stages (Studer, 1987; Warme, 1988), and constituted the High Atlas Belt. In the Central Jebilet, Triassic evaporite series outcrop a few kilometres South-East of Jbel Haïmer (Fig. 2a), and layers with sufficient thickness to be mined, about 100 km West of Jbel Haïmer in the Western Jebilet. The basement, now stripped of its Mesozoic cover as a result of erosion, was most probably continuously covered by Triassic evaporites during the Mesozoic.

Metal concentrations in brines: brines have high concentrations in Cu, Pb, Zn, (and Ag for Lwh) which were probably extracted from basement rocks (Fig. 16B). Among other potential source rocks present in the region, the black shales may release Ag, As, basic intrusions and some VMS lenses, Cu and Zn. These lithologies (metal sources) occur as weathered rocks just below the Triassic formations (brine source). Regoliths are generally considered as reserves of easily extractible metals by brines coming for upper levels (Boiron et al., 2010; Cathelineau and Boiron, 2019). Base metal contents in brines (Cu, Pb, Zn) are higher at Jbel Haïmer than at Roc Blanc (Table 6). Brines transported, therefore, specific metal contents depending on the deposit.

Mechanisms of Cu-ore deposition: metals might have been transported as chloride complexes owing to the high salinity of brines and the high Cu, Zn, and Pb of the Lwh brine (Yardley, 2005; Wilkinson et al., 2009). Thus, at temperatures below 300°C, chloride complexes are the predominant Cu, Pb, and Zn species, whereas hydrosulphides predominate at a higher temperature (> 400°C) in sulphide-saturated systems (Zhong et al., 2015).

Tm ice - Th pairs distribute between two end-members, the Lws and Lwh brines on the one hand, and a low salinity water end-member (Lw) on the other hand (Fig. 11). This trend corresponds to mixing between brine and a low-salinity fluid. The mixing is synchronous with the chalcopyrite deposition. The chlorinity decrease due to mixing may have destabilised metal (Zn, Cu, Pb) chloride complexes, and caused metal deposition (Wilkinson, 2001; Heijlen et al., 2003; Boiron et al., 2002, 2010; Neiva et al., 2015; Richard et al., 2016).

Moreover, Walter et al. (2019) emphasise on the example of the Schwarzwald district (SW Germany) that base metal (sulphide) ores form only if fluid mixing and sulfate reduction occur at the same time, TSR being the main process reducing sulfates from Triassic evaporites to produce H_2S. For those authors, a reducing agent (methane, for instance, or graphite in close host rocks) is required to precipitate ores. For example, Heijlen et al. (2008) noticed that the distribution of organic material determines the location of mineralisation in some parts of the Central African copper belt (DR Congo). At Jbel Haïmer, as at several studied areas in the Central Jebilet, graphite is abundant, and gas (CH_4, CO_2) storage in schist microporosity was documented (Bastoul, 1992). Thus, fluid mixing and reduction are the main factors controlling metal deposit at Jbel Haïmer. Likewise, Essarraj et al. (2017a) and Nshimiyimana et al. (2018) evoked a fluid reduction in graphite-rich black shale as a deposition mechanism for base metal and silver mineralisation at the Roc Blanc and Koudia El Hamra deposits, respectively.

Relative age of Cu-ore deposition: As Cu mineralisation at Jbel Haïmer occurs after the intrusion and deformation of the microdiorite dikes dated at 240 ± 10 Ma (Bouloton et al.,

2019), Triassic brine circulation during Mesozoic is, therefore, the most likely hypothesis. The thermal and tectonic stages related to the Central Atlantic Rifting are the main periods already proposed in a variety of other deposits for such brine migration (Jbel Ighoud barite deposit, Eastern Jebilet: Valenza et al., 2000; Roc Blanc Ag deposit: Essarraj et al., 2017a; Koudia El Hamra Ag deposit: Nshimiyimana et al., 2018). Furthermore, large basaltic flows, coeval with the Central Atlantic Ocean opening are reported in the region (Frizon de Lamotte et al., 2008, 2009). They are witnesses of abnormal thermal gradients recorded in the Western Meseta between 250 and 210 Ma (Ghorbal et al., 2008; Saddiqi et al., 2009; Barbero et al., 2011), and around 195 ± 4 Ma (Huon et al., 1993).

At Jbel Haïmer, tectonic traps consist of clusters of reactivated and re-opened Variscan (N-S to NE-SW and NW-SE) discontinuities during post-Variscan events. The nearby presence of the MSZ and the Aït Bella OuSaïd shear zone may constitute drainage zones for brines (Fig. 2a and 16B), as they are reactivated during the Central Atlantic rifting (Laville and Piqué, 1991; Piqué and Laville, 1993, 1996).

Similar ore deposition processes in Morocco: Early fluid stages synchronous with the Panafrican magmatism and metamorphism, followed by later migration of evaporitic brines during the Mesozoic are typical of most Anti-Atlas deposits: Zgounder (Ag-Hg), Bou Azzer (Co-Ni-Ag-Au), and Imiter (Ag-Hg) deposits (Essarraj et al., 1998, 2005, 2016, 2017b). Similarly, fluid flows induced by extensional tectonics, contemporary with the first stages of the Atlantic Ocean opening, are proposed for the Assif El Mal deposit, High Moulouya district and Tirhza-Jbel Aouam district (base metal – polymetallic, Morocco) (Bouabdellah et al., 2009; Margoum et al., 2015; Rossi et al., 2016 respectively).

Conclusion

590 This study concludes a general framework of fluid circulations and mineral deposition at the
591 Jbel Haïmer area summarised as follows:

592 - The first stage is related to fluid convection induced by Late Variscan magmatism.
593 High-temperature fluids having a typical metamorphic signature (C-O-N-H) deposited
594 minor concentrations of Sn-As-(Co-Ni) in quartz veins.
595 - The second stage occurs much later and is independent of the first stage. It is
596 characterised by the introduction of metals (Cu – (Zn – Pb), Ag) in the system by
597 primary evaporitic brines characterised by their salinity and enrichment in K, Mg. The
598 likely origin of the brines is the Triassic evaporitic levels deposited above the basement
599 unconformity. The high metal concentrations in fluid inclusions are considered as an
600 excellent pathfinder of exploration targets. The high chlorinity of the brines favoured
601 the extraction of metals from basement rocks, especially those affected by weathering
602 (regolith) located just below the unconformity and Triassic formations.
603 - Cu-ores deposited at medium temperatures (220-280°C) corresponding to a burial of
604 around 4-5 km. The main driving force for Cu-deposition is the dilution of ore-forming
605 brines by low salinity fluids, probably of meteoric origin.
606 - Tectonic traps are issued in part from Variscan deformation which developed damaged
607 zone along shear zones. Mesozoïc extensional tectonics reactivated inherited shear
608 zones during post-240 Ma events. Although the absolute age of the Cu-ore stage is not
609 available, it is likely related to the Atlantic Ocean rifting stages. The latter favoured the
610 downward migration of Triassic brines into the basement. Such processes are very
611 similar to those proposed for the Anti-Atlas, High Atlas and Meseta silver and base
612 metal deposits.
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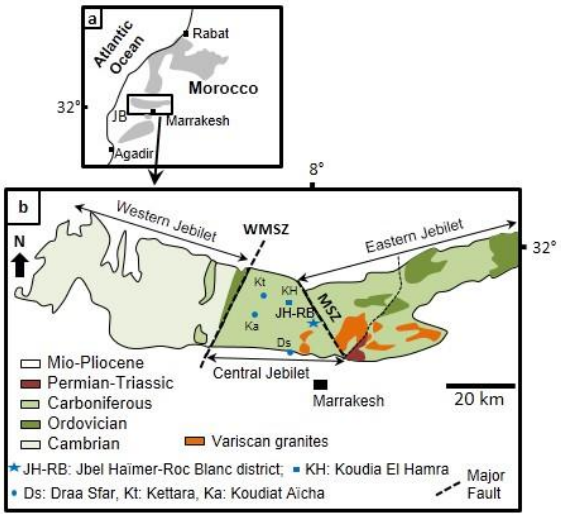


Fig. 1. a Location of the Variscan massifs in Morocco (in grey: central-meridional domain), JB: Jebilet. **b** Geology of the Jebilet massif (Huvelin, 1977; in Bernard et al., 1988), and location of the main deposits: the Jbel Haïmer and Roc Blanc (Ag) district, the Koudia El Hamra Ag deposit, and the main volcanogenic massive sulphide deposits (VMS pyrrhotite-polymetallic lenses): Draa Sfar, Koudiat Aicha and Kettara, from Huvelin (1977). (MSZ: Marrakesh Shear Zone; WMSZ: West Moroccan Shear Zone, see text for references).

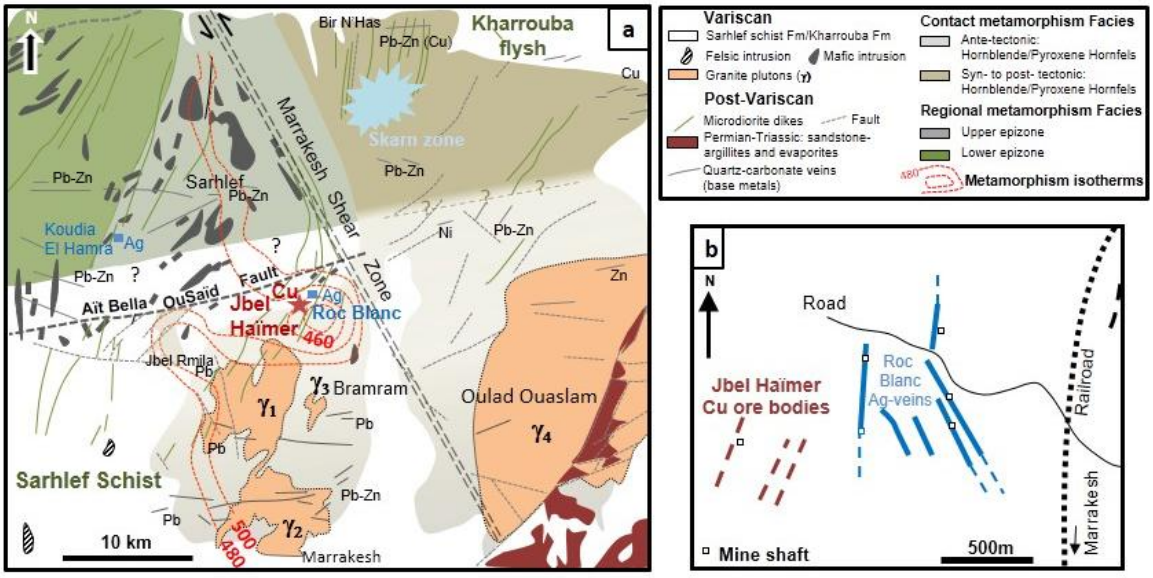
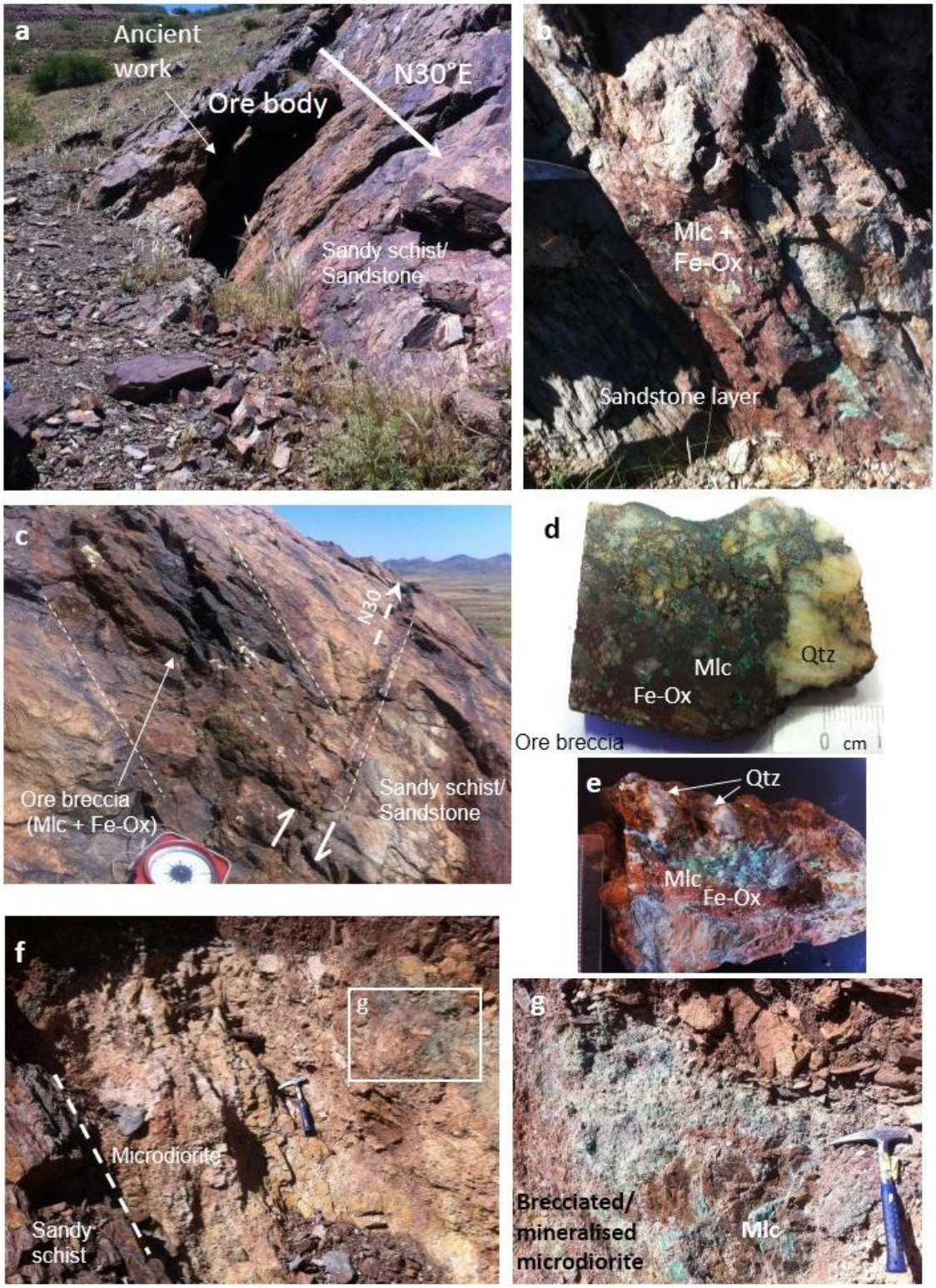


Fig. 2. a Location of base metal and silver ore veins (deposits and occurrences) with the indication of major faults in the Central Jebilet around Jbel Haïmer and Western part of the

Eastern Jebilet (from Huvelin, 1972); only kilometre size faults are represented. γ: granites 1, 2, 3 and 4: Tabouchennt, Bamega, Bramram and Oulad Ouaslam respectively. The Jbel Haïmer–Roc Blanc district is detailed in b. Metamorphism facies and isotherms of thermal metamorphism M2 are reported from Delchini et al. (2018) and Delchini (2018); skarn zone: from Huvelin (1977), studied by Bastoul, (1992). **b** Schematic map of the mining works of the Jbel Haïmer (main ore bodies) and Roc Blanc Ag- ore bodies (From Huvelin 1977).



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954 **Fig. 3. a**, **b**, **c**, **f**, **g** Field exposures of the Jbel Haïmer area showing gossan: oxidised ores
955 (malachite (Mlc) and Fe-oxides (Fe-Ox) sandy schist/sandstone and brecciated/altered
956 microdiorite dike. **d** and **e**: hand specimens showing ore breccia with quartz (Qtz) fragments
957 impregnated with malachite and Fe-oxides

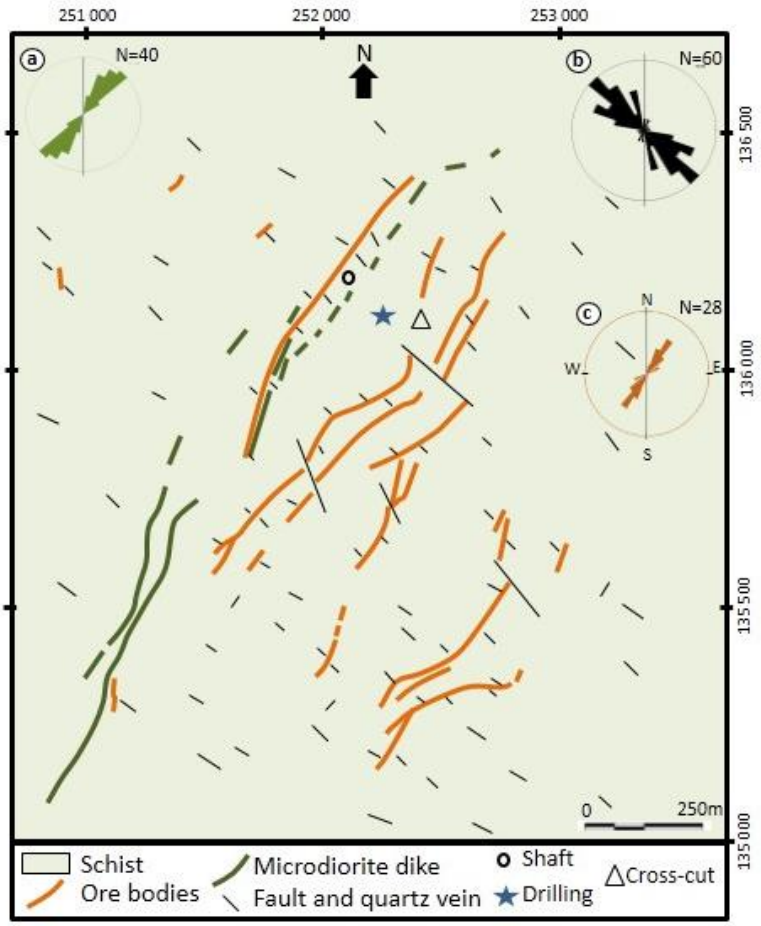
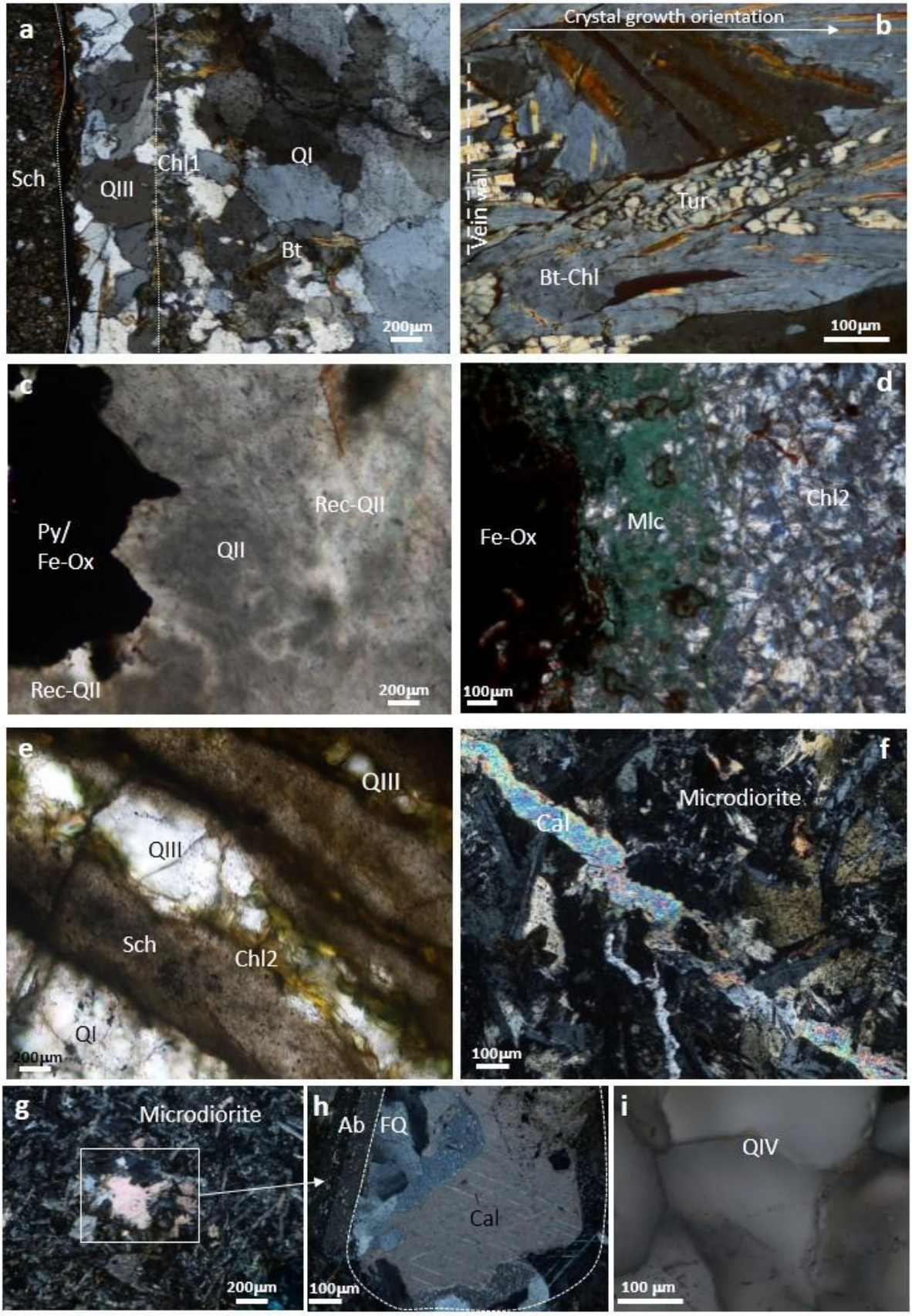


Fig. 4. Representation of fractures, main ore bodies, and microdiorite dikes subparallel to local schistosity (S1) in the Jbel Haïmer area. **a**, **b**, **c** rose-diagrams corresponding respectively to schistosity, fractures, and veins, and ore bodies (Boutira and Hakkou, 2012; in Essarraj et al., 2013), N: number of measurements.

Fig. 5. a, **b**, **c**, **d** Relationships between schistosity, fractures and veins crosscutting sandy schist and sandstone layers; QI and QII: early quartz (see text for description). **e** fractures in microdiorite, **f** Calcite veinlet crosscutting microdiorite, **g** Feathery quartz veinlets crosscutting microdiorite.



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969 **Fig. 6.** Mineralogy of Jbel Haïmer. **a** NW-SE vein with deformed and cloudy quartz (QI) and

970 chloritised biotite (Chl1)-(Bt), crosscutting schist (Sch), and newly formed quartz veinlet with

undeformed quartz (QIII). **b** chloritised biotite (Bt-Chl) and brecciated tourmaline crystals (Tur) perpendicular to the QI vein wall. **c** Cloudy QII quartz with strongly oxidised pyrite (Py/Fe-Ox) in the core of QII vein, Rec-QII: recrystallisation microdomains in QII around pyrite. **d** Malachite (Mlc), Fe-oxides (Fe-Ox), and chlorite 2 (Chl2) invading the QI vein core. **e** Undeformed QIII-chlorite 2 (Chl2) veinlet parallel to QI (deformed) vein, both crosscutting schist (Sch). **f** calcite veinlet crosscutting microdiorite. **g** Micro-vacuole in microdiorite filled with late minerals (detail in H). **h** Undeformed feathery quartz (FQ), then calcite (Cal) filling a vacuole in microdiorite, Ab: albite from microdiorite. **i** Late geodic quartz QIV poor in fluid inclusions and FIP.

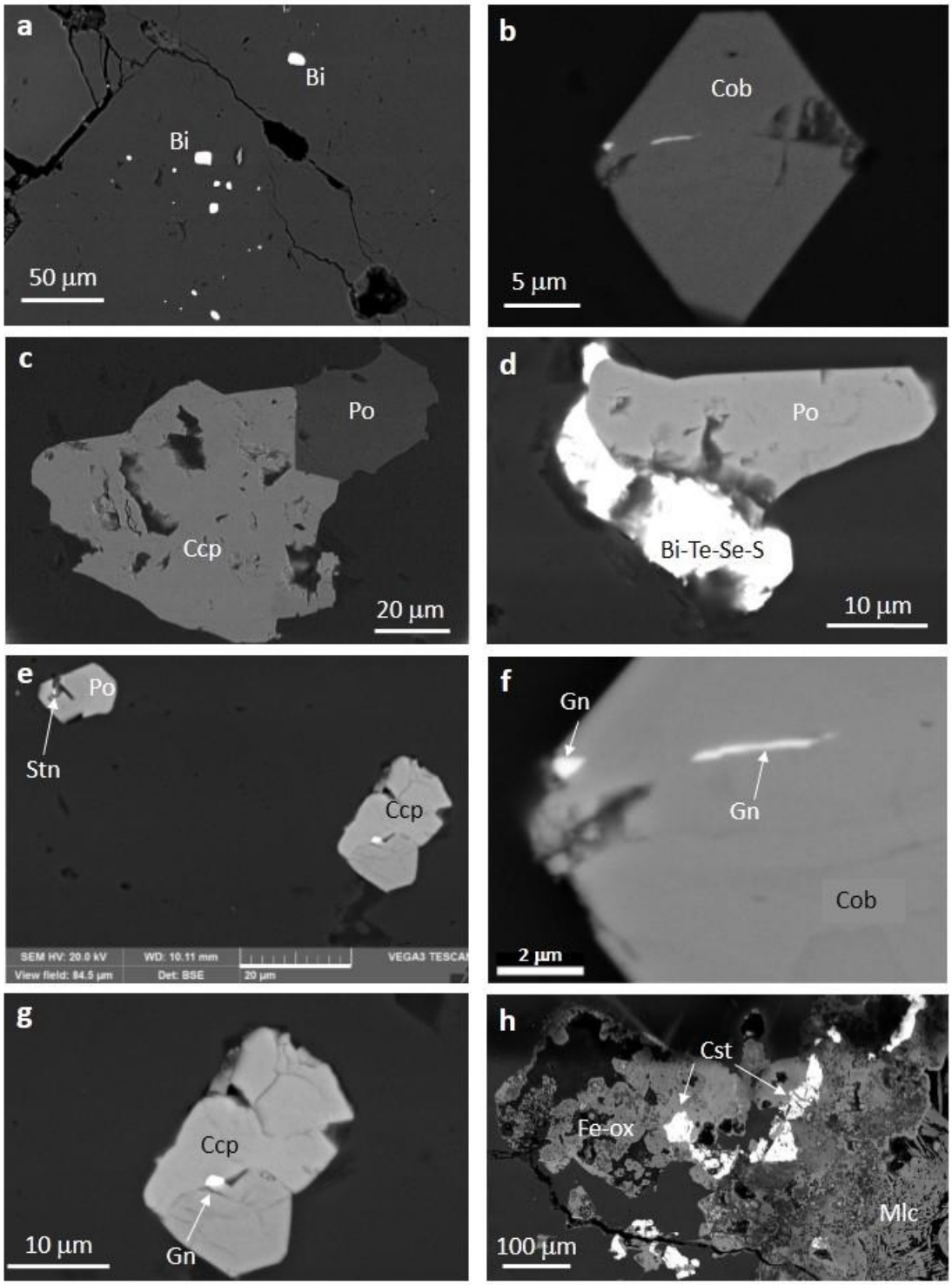


Fig. 7. Ore minerals from Jbel Haïmer observed by scanning electron microscope. **a** Bismuth and bismuthinite blebs (Bi). **b** cobaltite euhedral crystal (Cob). **c** Chalcopyrite (Ccp) crystallising over pyrrhotite (Po). **d** Bi-Te-Se-S mineral growing over pyrrhotite (Po). **e** Stannite (Stn) in fracture of pyrrhotite (Po). **f** galena (Gn) in microfracture crosscutting cobaltite (Cob). **g** Galena (Gn) in microfracture crosscutting chalcopyrite (Ccp). **h** Malachite (Mlc), Cassiterite (Cst), and Fe-oxides in oxidised ore.

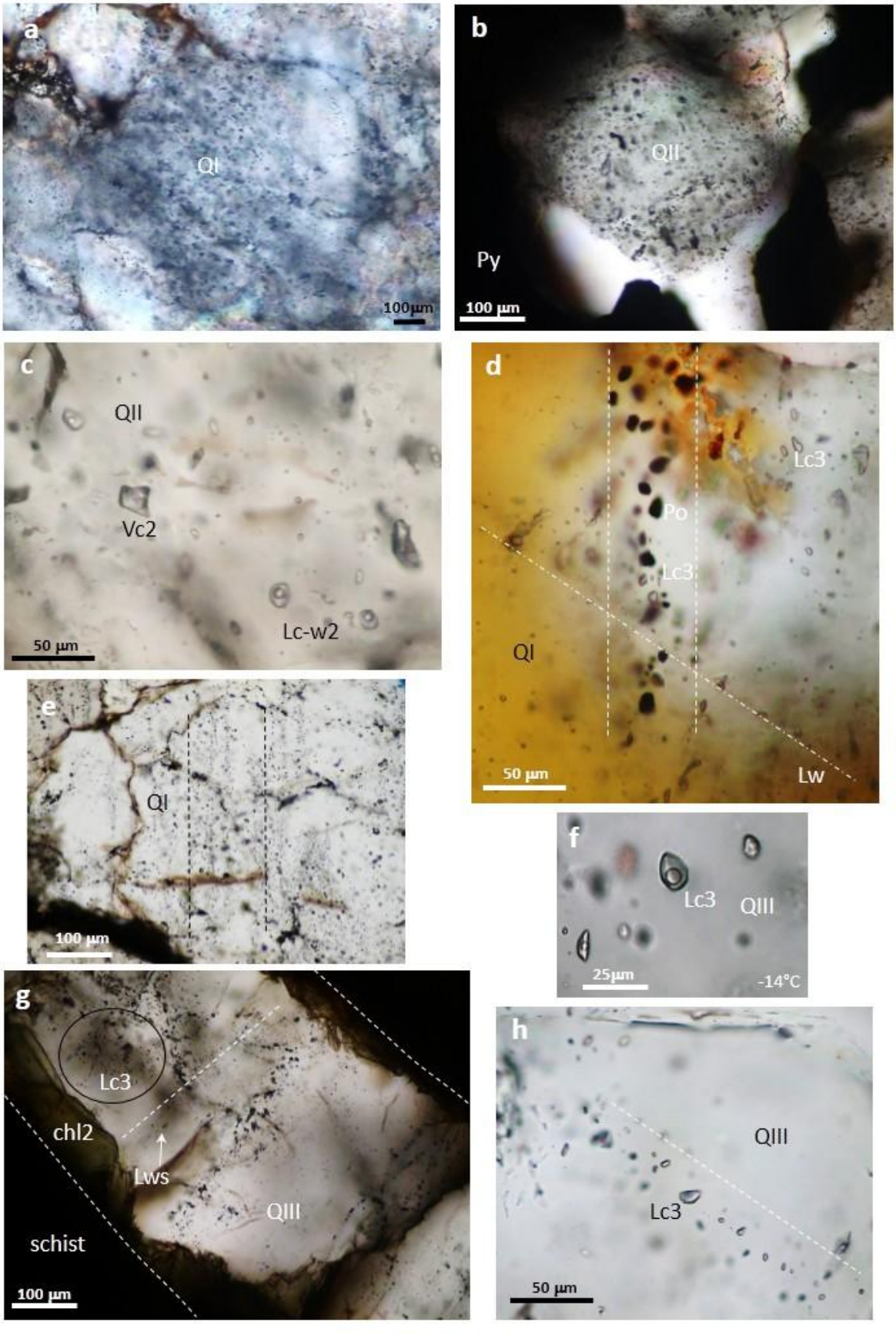

a
QI
100µm
b
QII
Py
100 µm
c
QII
Vc2
Lc-w2
50 µm
d
Lc3
Po
Lc3
QI
50 µm
Lw
e
QI
100 µm
f
Lc3
QIII
25µm
-14°C
g
Lc3
chl2
Lws
QIII
schist
100 µm
h
QIII
Lc3
50 µm

Fig. 8. Fluid inclusions (FIs) from Jbel Haïmer: early stages. **a** QI darkened by FIs clouds from different generations. **b** QII quartz recrystallised around pyrite crystals (Py) and late sulphides assemblages. **c** FIs in QII. **d** Lc3 FIs synchronous with small pyrrhotite crystals: Lc3 FIs and pyrrhotite microcrystals are trapped in the same microfracture crosscut by an aqueous fluid inclusion (Lw) plane. **e** Lc3 FIP network parallel to NW-SE QI veins. **f** Lc3 FIs (two-phase at -14°C). **g** Lc3 FIs cloud (circle) in NW-SE QIII – chlorite 2 (Chl2) vein, and Lws aqueous FIP perpendicular to QIII vein. **h** Lc3 FIs pseudo-secondary in QIII vein. See text for description of fluid inclusion types.

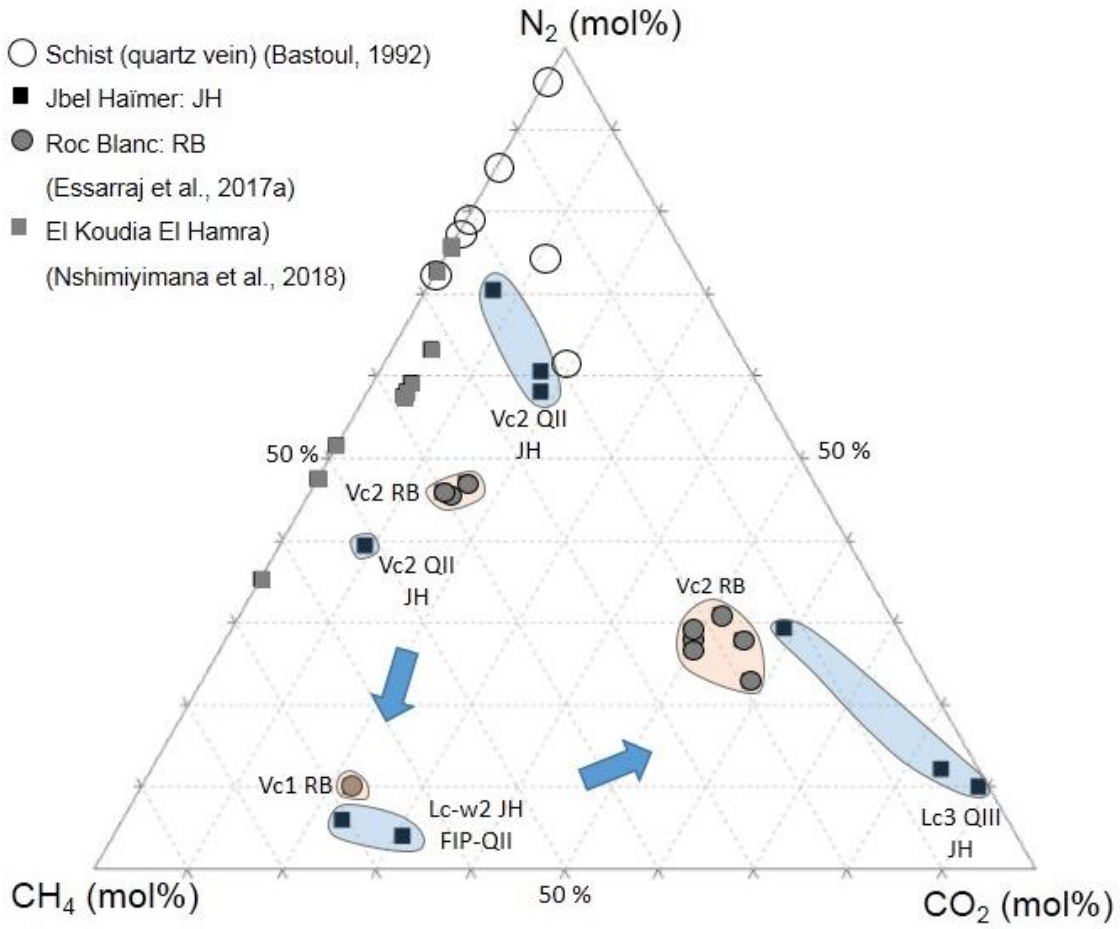


Fig. 9. CO_2-CH_4-N_2 ternary diagram presenting Raman data for FIs from Jbel Haïmer compared with data from vapour fluid inclusions from the Roc Blanc and the Koudia El Hamra silver deposits and a quartz vein (Bastoul, 1992) crosscutting the Sarhlef schist in central Jebilet. Arrows indicate the evolution of fluid inclusion composition over time.

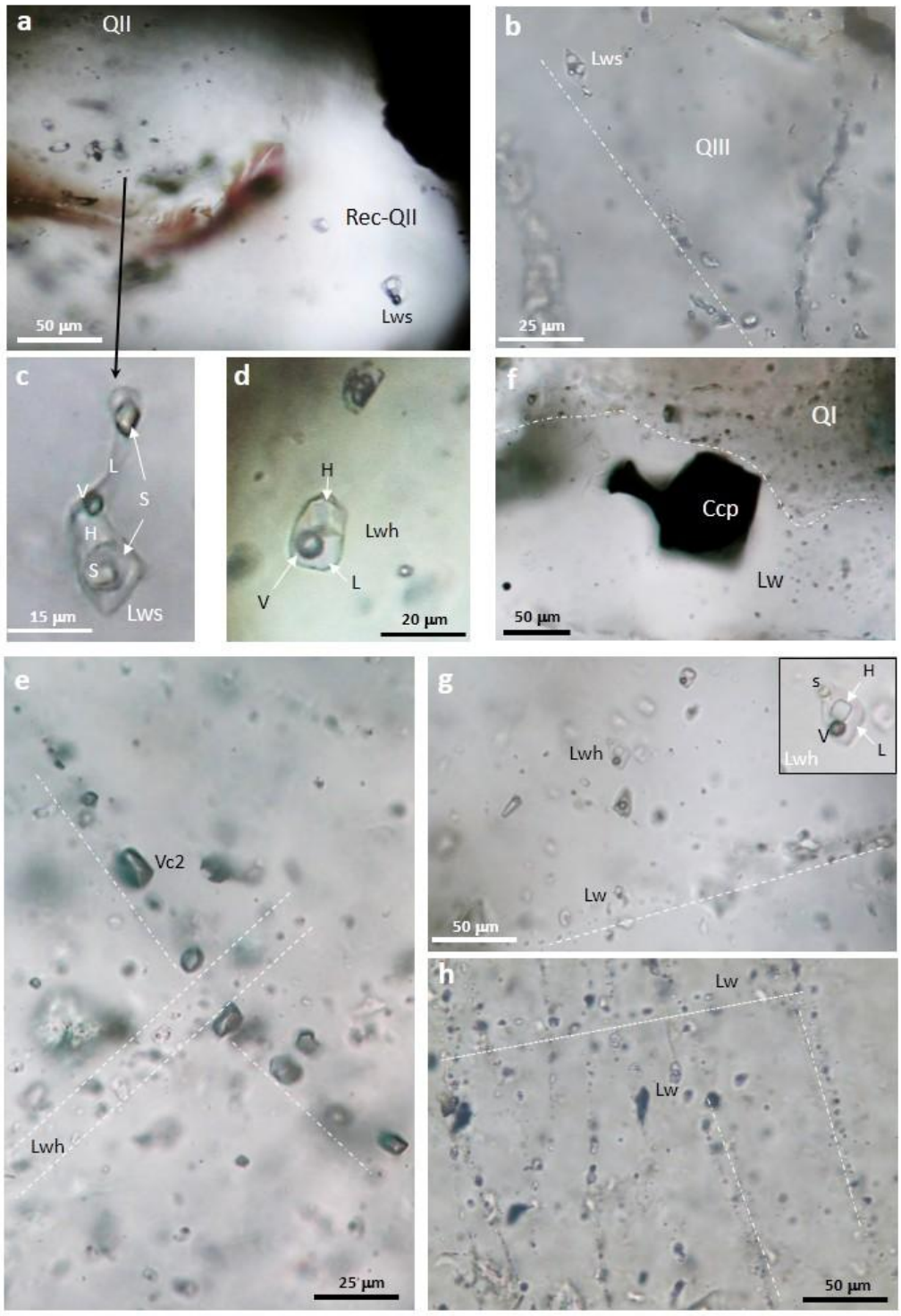


Fig. 10. Fluid inclusions from the late stages at Jbel Haïmer. **a** Lws FIs in recrystallised QII quartz (Rec-QII). **b** Lws FIPs in QIII. **c** Lws FI showing halite cube and 3 solids. **d** Lwh FI. **e** Lwh FIs crosscutting one phase carbonic FIP in QIII. **f** recrystallised cloudy quartz QI around chalcopyrite crystal (fluid inclusion poor), scarce Lw FIs. **g** Aqueous FIP (Lw) crosscutting Lwh FI assemblage in QIV. **h** Perpendicular FIP sets crosscutting QI quartz vein.

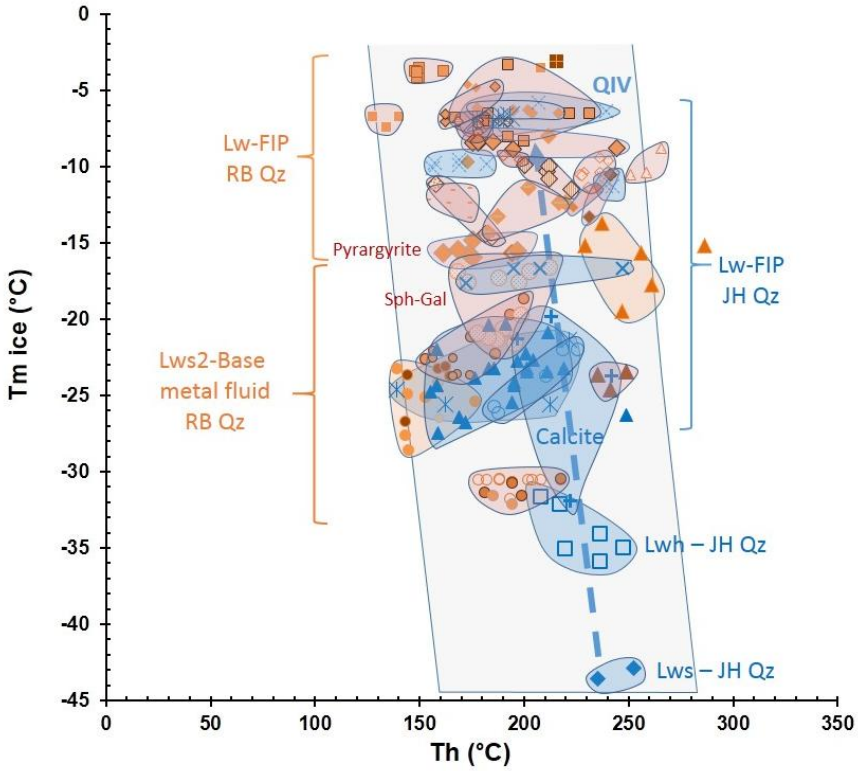


Fig. 11. Tm ice vs Th for aqueous fluid inclusions from Jbel Haïmer (in blue). For comparison, data for aqueous fluid inclusions from the Roc Blanc deposit are presented (in orange) (from Essarraj et al., 2017a). Different symbols correspond to different FIPs. Arrow represents the mixing trend described in the text. JH Qz: FIs in quartz from Jbel Haïmer, RB Qz: FIs in quartz from Roc Blanc, calcite domain: FIs in calcite from Jbel Haïmer. Sph-Gal and pyrargyrite: FIs associated respectively to sphalerite – galena and pyrargyrite from Roc Blanc (Essarraj et al., 2017a). See text for fluid inclusions types and comments.

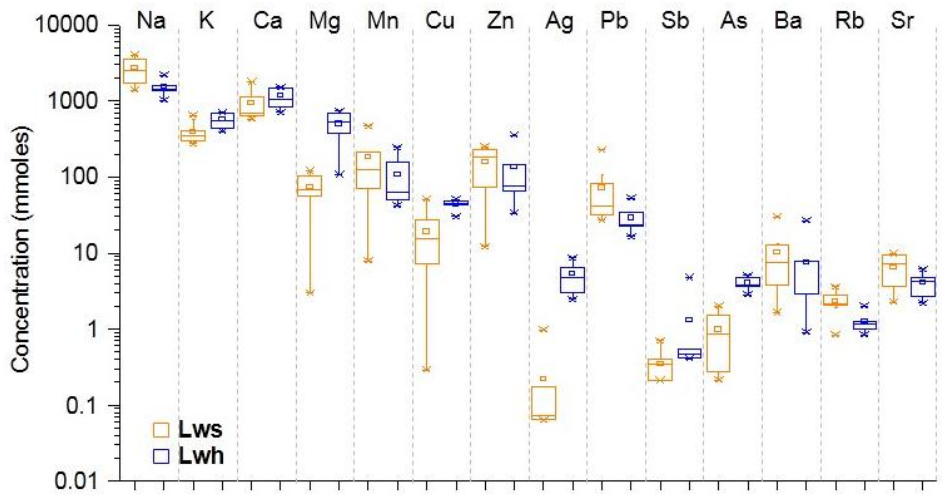


Fig. 12. Range of the major, minor, and trace element compositions (in mmol/kg of solution) of the Lws and Lwh brines from Jbel Haïmer, analysed by LA-ICP-MS.

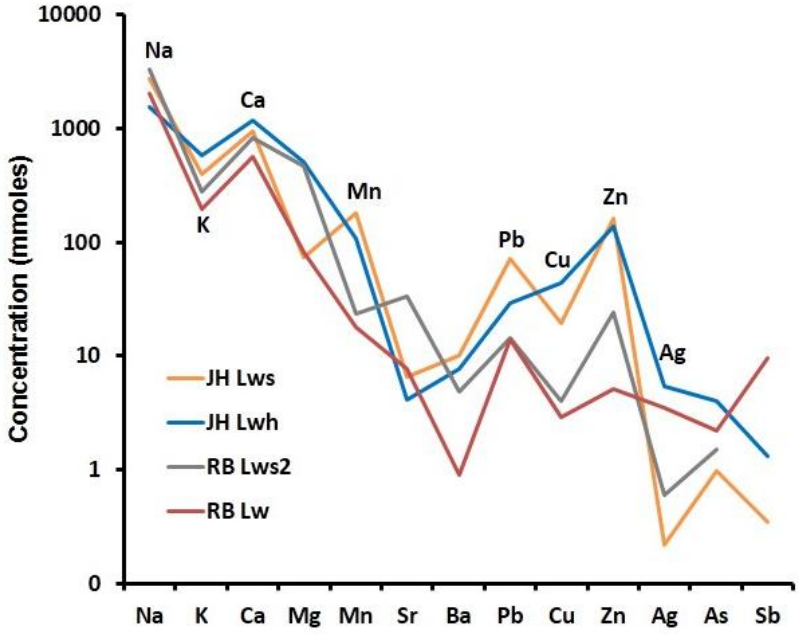


Fig. 13. Comparison of the Jbel Haïmer brine compositions (JH Lws and JH Lwh brines) with the compositions of brines from the Roc Blanc silver deposit (RB Lws2 and RB Lw brines, from Essarraj et al., 2017a).

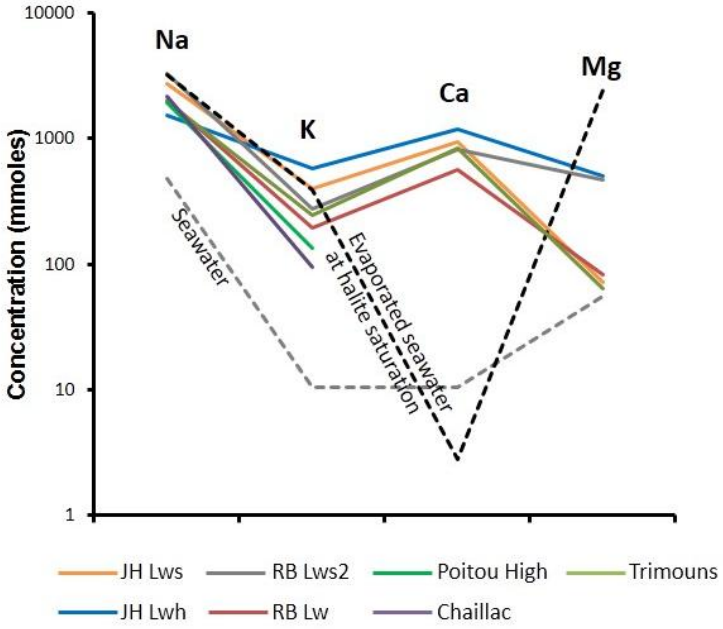
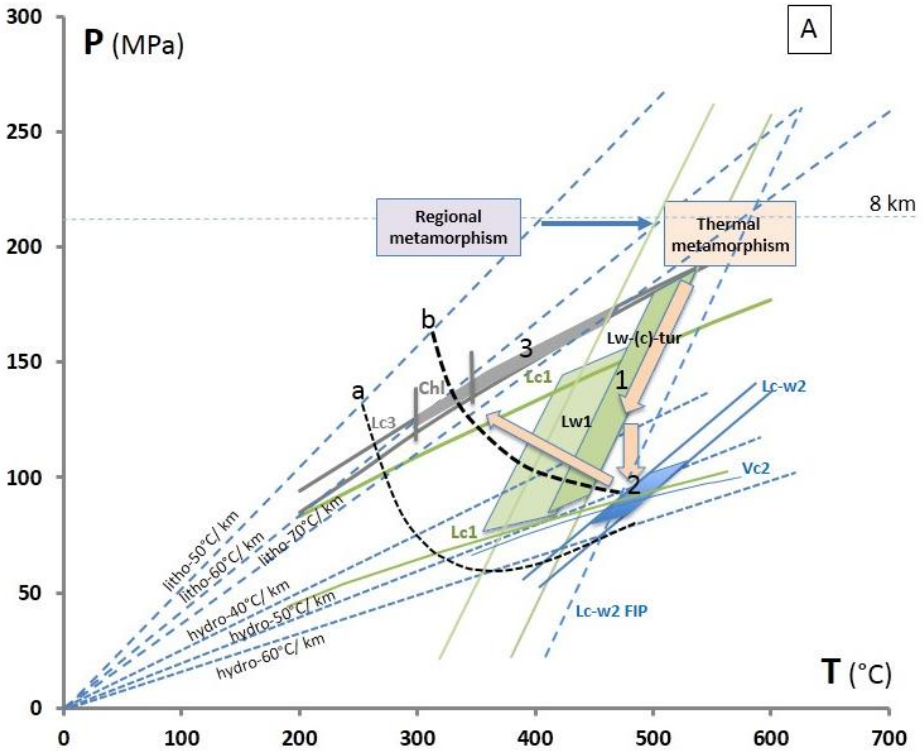


Fig. 14. Average Na, K, Ca and Mg concentrations in fluids from Jbel Haïmer (JH Lws and JH Lwh brines, this study) and Roc Blanc (RB Lws2 and RB Lw brines, Essarraj et al, 2017a). For comparison, chemical composition of brines from literature are plotted: Poitou High and Chaillac (Boiron et al., 2002), Trimouns (Leisen et al., 2012b) as well seawater and evaporated seawater after halite saturation (Fontes and Matray, 1993).



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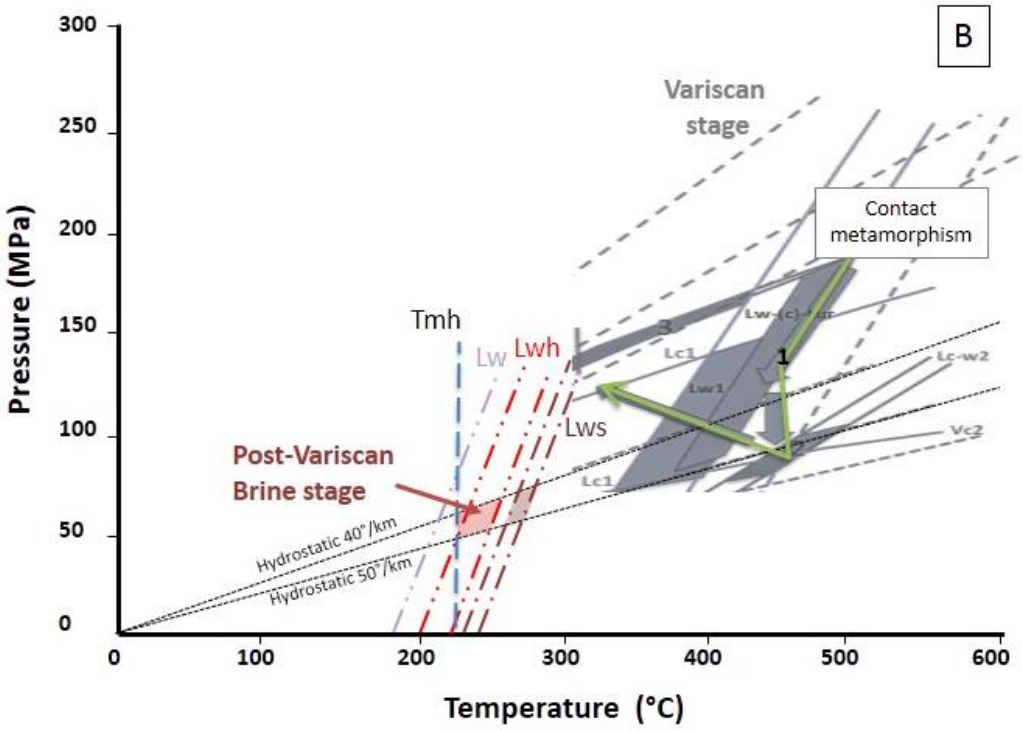

1035

1036 **Fig. 15.** P-T diagram for the two main stages at Jbel Haïmer with typical inclusions of each
1037 fluid inclusion type, distinguished by their colours: **A** Variscan stages; **B** Post-Variscan Cu-
1038 base metal ore stage. Thermal gradients for the lithostatic and hydrostatic pressures are
1039 reported, as well as isopleths from Weisbrod (1984) for two compositions: (a) 10 % CO_2 and 6

% NaCl and (b) 20 % CO_2 and 6 % NaCl. The thermal metamorphism box for the Central Jebilet is from Bastoul (1992). Chl: Roc Blanc chlorite temperature domain from Essarraj et al. (2017a). Arrows represent the evolution P-T trend. In Fig. 15B, the Variscan domain is shown in grey.

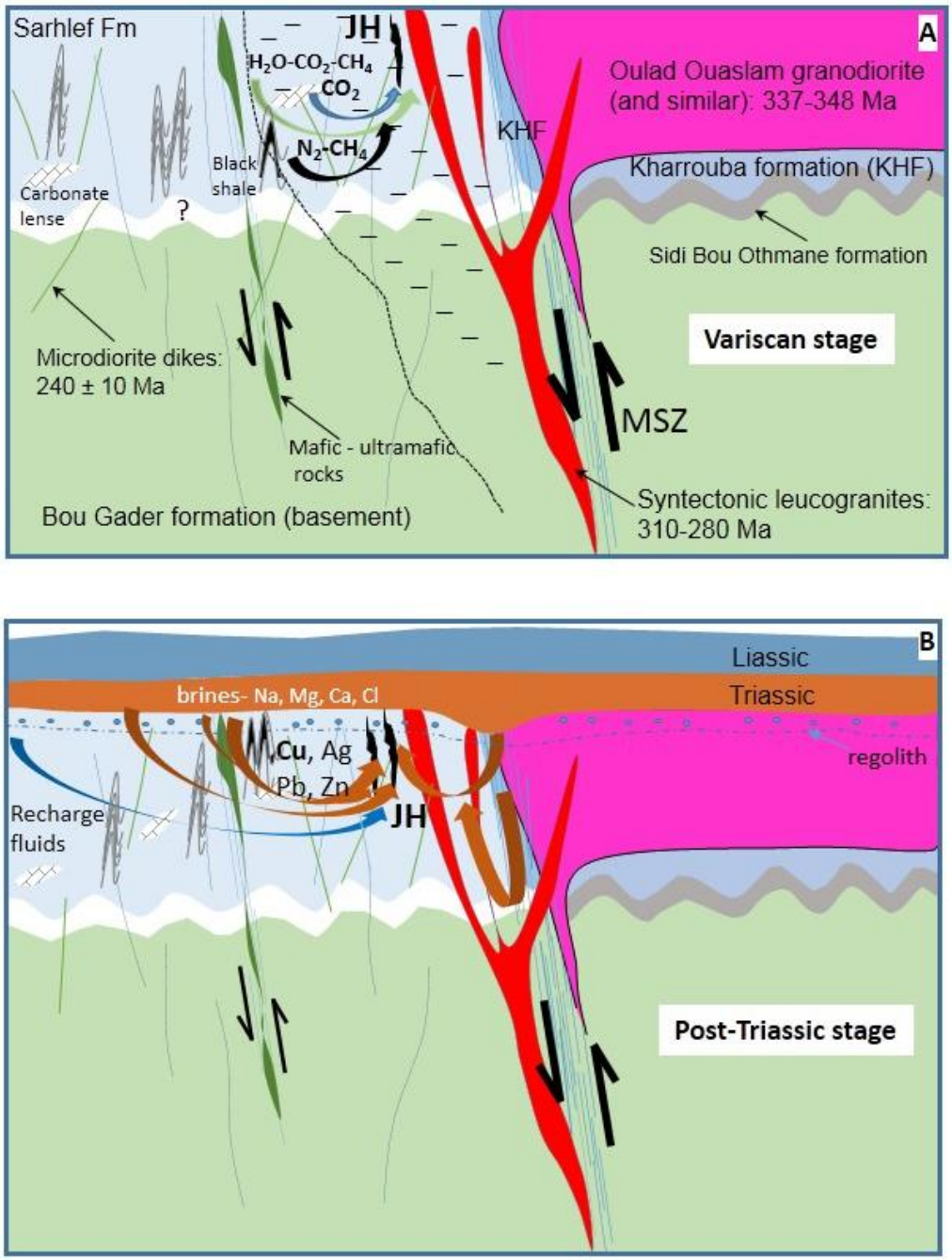

Fig. 16. A two-stage model of fluid circulation in the Jbel Haïmer area (JH) Jebilet: **A** Variscan stage: the intrusion of late leucogranites, the thermal metamorphism M2a lead to the genesis of CO_2, CH_4-N_2 and H_2O-CH_4-CO_2 fluids which mix, and form, in pre-existing discontinuities,

the early quartz assemblages with the Fe-As-Sn-(Bi) association. See text for references on the granodiorite, granite and microdiorite ages. MSZ: Marrakesh Sear Zone. **B** Mesozoic brine stage: after peneplaination and deposition of Triassic/Early Jurassic formations, circulation of brines coming from Triassic evaporite formation within the regolith, and several types of lithologies including black shales, basic intrusions and schists, yield to the extraction of metals, the transport and the formation of Cu-polymetallic ores in the Jbel Haïmer area. Triassic and Liassic covering rocks in the Jebilet are from Saddiqi et al. (2009). Figures (A) and (B) are not to scale.

Table 1. Paragenetic sequence for the Jbel Haïmer occurrences.

Stages / Minerals	Pre ore stages			Ore stages	
	Sn-Bi stage	Fe-As-Co-Ni stage	Fe stage	Cu-Ag-(Au)- base metals stages	Supergene
Quartz I (QI)	▬				
Biotite (chloritized: chlorite 1)	▬				
Turmaline	—				
Muscovite	? ↕ —				
Cassiterite	—				
Bi-minerals	—				
Quartz (QII)		▬			
Pyrite		▬			
Arsenopyrite		—			
Co-pyrite		—			
Co-Fe-Ni-sullfarsenide		—			
Quartz III (QIII)			▬		
Pyrrhotite			—		
Chlorite 2			▬		
Feathery quartz				—	
Carbonates				▬	
QIV				—	
Fe-Sphalerite				—	
Chalcopyrite				? ▬	
Stannite				—	
Bi-Te-Se sulphide				—	
Galena				▬	
Natives : Ag-Au alloy				—	
Malachite					▬
Covellite					—
Fe-Oxides/hydroxides					▬
Cerussite					—
Native metals					—
(Pb, Zn, Ag, Cu, Bi, Se)					
Fluid circulations	N2-CH4-CO2-H2O metamorphic fluids - Variscan			Basinal brines (aqueous) - Triassic +	Meteoric Fluids

Brecciation

Table 2. Description of fluid inclusion types from Jbel Haïmer using nomenclature from Boiron et al. (1992) completed. L: liquid, V: vapour; (*): fluid inclusions homogenising to an aqueous liquid.

Fluid inclusion type	Fluid inclusion description
Lc	One-phase liquid N2-CH4-CO2 (carbonic: c)
Vc	One-phase vapor N2-CH4-CO2 (carbonic: c)
Lc-w *	Two-phase aqueous (w) carbonic (c), liquid water dominated + V
Lw-(c) *	Two-phase liquid water dominated + V, with traces of gas (clathrates)
Lws *	Multi-phase aqueous (w) L+V + unidentified solids + halite
Lwh *	Multi-phase aqueous (w) L+V + one solid + halite
Lw *	Two-phase liquid water dominated + V

Table 3. Summary of microthermometry data of fluid inclusions from the Jbel Haïmer ores. Tm CO_2: melting temperature of volatile phase, Th CO_2: homogenisation temperature of volatile phase to the liquid (L) or vapour (V), Tm cl: melting temperature of clathrate, Te: eutectic temperature, Tm ice: melting temperature of ice, Th: homogenisation temperature to the liquid (L) or vapour phase (V), Tm hh: melting temperature of hydrohalite, Tm h: melting temperature of halite. All values are in °C, mode: bold font in brackets

Fluid inclusion types			**Microthermometry data**			
Early stages - Fe-Sn-As-Bi-Co-Ni	Tm CO2 (°C)	Th CO2 (°C)	Tm cl (°C)	Te (°C)	Tm ice (°C)	Th (°C)
Aqueous dominated FI in quartzite						
Lw-(c)-tur - Scattered in tourmaline (< 10 µm, 60-70% liq. H2O; 3 FIs)			~ -0.6 / 0	-52 / -50	-25.7 / -22.3	> 380 ± 10 L
Lw-(c)1 - Secondary in quartz (< 10 µm 60-70 % liq. H2O; 11 FIs)			-1 / 0	-51 / -49	-12.6 / -11.7	286 to > 386 L
Fluid inclusions in QI						
Lw1 - Random (10 µm, 60-70% liq. H2O; 5 FIs)			Tm hh: ~ -13	-50 / -49	-25 / -16.5	296 to > 357 L
Lc1 - Random (< 15 µm; 20 FIs)	-78.4 / -67.9 L **(-71)**	-70.2 / -13.7 L **(-32)**				
FIP Lw2 - (< 10 µm, 70-90% liq. H2O; 25 FIs)				-50 / -44	-11.8 / -8.1 **(-10)**	261 - 317 **(290)**
Fluid inclusions in QII (< 20 µm): Vc2 and Lc-w2						
Scattered: Vc2 (0 to 10% H2O) and Lc-w2 (up to 70% liq. H2O; 15 FIs)	-102.8 / -68.5	-71.6 / -60.5 V	-4.2 / 1 **(-3)**		-8.4 / -3.8 **(-4)**	> 460 L ; D > 350
FIP: Vc2 (0 % visible H2O) and Lc-w2 (up to 70% H2O; 11 FIs)	~ -102	~ -97 V	6.2 / 9.1		-10 / -9	414 - 453 L **(425)**
Fluid inclusions in QIII						
Lc3 - Scattered to pseudo-secondary (65 FIs) (< 20 µm)	-65 / -57.6 **(-59)**	-46.9 / +18.6 (L, V)				
Ore stages - Cu-Zn-Pb-Ag-(Au) - Aqueous fluids		Tm h (°C)	Tm hh (°C)	Te (°C)	Tm ice (°C)	Th L-V (°C)
Lws (< 20 µm; 20 - 30 % vap; 6 FIs)		209 to > D ~ 230	-28 / -15.9	-65 / -50	-43.6 / -39	174 to 252 - D ~ 230 ° and >
Lwh (< 15 µm, 10% vap; 43 FIs)		155 / 262 **(220)**	~ -12 / -13	-65 / -50	-35.9 / -31.6	179 / 264 L **(220) ;** D ~ 270
Lw / Feathery quartz - calcite (scattered and FIP; 9 FIs)					-32 / -20	186 - 242 L
Lw FIP/QI, QII, QIII (< 15 µm, 10% vap; ~ 50 FIs)			~ -11.9 / -6.5	-60 / -45	-29.6 / -7.1	139 - 249 L (**200**)
Lw FIP/QIV (10 µm, 10% vap; 10 FIs)				~ -49	-11.3 / -5.8	156 / 242 L

Table 4. Microthermometric and Raman data for carbonic and aqueous-carbonic fluid inclusions from Jbel Haïmer and the neighbouring Roc Blanc deposit (*Essarraj et al., 2017a). Concentrations of the volatile phase components are given in mole %. Vc1 and Vc2 are respectively early and late carbonic vapour fluid inclusions from Roc Blanc. (FIs: Fluid inclusion, nm: not measured)

Deposit/Fluid inclusion type	FI reference	CO2	CH4	N2	Tf CO2	Th CO2
Jbel Haïmer						
Vc2 scattered/QII	L6-2A-P2-8	7.7	22.5	69.8	-102.8	(nm) V
Vc2 (~10% H2O) scattered/QII	L6-2C-IF4	17.7	22.1	60.2	-73.6	~ -70
Vc2 scattered/QII	L6-2C-IF5	18.6	23.8	57.6	-82.1	(nm) V
Vc2 scattered/QII	L6-2A-P2-9	9.1	51.5	39.4	-95.6	(nm) V
Lc-w2 scattered/recrystallised QII	L6-2A-P1-1	30.7	64.5	4.8	~ -102	-96.7 V
Lc-w2 scattered/recrystallised QII	L6-2A-P1-2	23.8	70	6.2	(nm)	(nm) V
Lc3 scattered QIII	JH4-1-IF19	58.5	12.2	29.3	-59.6	-13.5 L
Lc3 peudo-secondary QIII	JH4-1-IF1	88	1.7	10.3	-59.5	-15.9 L
Lc3 pseudosecondary QIII	JH4-1-IF5	83.7	4	12.3	-58.2	+1.6 L
Roc Blanc *						
Vc1	RB19/2	22.5	67.3	10.2	(nm)	(nm) V
Vc1	RB22/1-1	15.9	36.9	47.3	-81.1	(nm) V
Vc1	RB22/1-2	14.2	39.6	46.2	-79.1	-106.8 V
Vc1	RB22/1-6	14.8	39.6	45.5	-80.6	(nm) V
Vc2	RB22/1-3	50.2	22.5	27.3	-68.7	-51.8 V
Vc2	RB22/1-4	48.9	22	29.1	-69	-51.3 V
Vc2	RB19B-1	58.2	19.1	22.8	-60.8	-23.6 V
Vc2	RB19B-3	54	17.5	28.5	-60.7	-24.7 V
Vc2	RB19B-5	51.2	17.7	31.1	-59.2	(nm) V

Table 5. Representative LA-ICP-MS data obtained on Lws and Lwh fluid inclusions from the Cu- stage (Jbel Haïmer). Absolute element concentrations are in mmol/kg of solution in the first line, and ppm (second line, italic font).

FI type	FI number	Na	K	Ca	Mg	Mn	Cu	Zn	Ag	Pb	Sb	As	Ba	Rb	Sr
Lws	**JH3-6-1-IF1**	2971	277	860	61	146	13.9	251.2	0.1	68.6	0.7	1.1	12.2	3.6	9.8
		68332	*10814*	*34405*	*1452*	*8027*	*878*	*16577*	*7*	*14194*	*85*	*79*	*1698*	*310*	*861*
	JH3-6-1-IF3	2284	649	601	97	183	4.9	244.9	0.2	228	0.3	0.9	4.5	3.1	4.3
		52539	*25300*	*24034*	*2330*	*10072*	*311*	*16163*	*19*	*47196*	*42*	*64*	*623*	*265*	*377*
	JH3-6-2-IF2	2542	415	1304	56	98	15.4	72.8	nd	32.2	0.2	2.0	3.7	0.8	3.7
		58473	*16182*	*52148*	*1354*	*5401*	*973*	*4807*	*nd*	*6658*	*26*	*149*	*519*	*72*	*322*
	JH3-6-3-IF1	1753	388	1079	76	431	50.6	226.8	0.1	81.5	0.4	0.3	12.2	2.1	8.1
		40317	*15126*	*43163*	*1817*	*23704*	*3185*	*14968*	*7*	*16870*	*50*	*21*	*1701*	*175*	*707*
	JH3-6-4-IF4	3977	328	580	30	66	31.2	147.3	0.1	39.1	0.2	0.2	7.5	2.1	3.4
		91475	*12806*	*23199*	*730*	*3655*	*1965*	*9720*	*14*	*8083*	*26*	*16*	*1042*	*182*	*299*
	JH3-6-4-IF5	3539	354	706	3	125	0.3	213.0	nd	41	0.2	nd	12.5	2.8	7.3
		81400	*13791*	*28238*	*71*	*6858*	*18*	*14058*	*nd*	*8489*	*26*	*nd*	*1742*	*238*	*637*
	JH3-6-7-IF1	1621	304	1139	69	471	17.4	212.5	1.0	108.0	0.4	0.7	29.8	1.9	9.3
		37291	*11875*	*45570*	*1660*	*25912*	*1099*	*14028*	*103*	*22349*	*54*	*55*	*4145*	*163*	*812*
	JH3-6-7-IF4	3245	376	632	121	212	7.3	180.2	0.1	68.3	0.4	0.3	13.2	2.6	7.4
		74646	*14652*	*25284*	*2893*	*11674*	*462*	*11896*	*8*	*14143*	*53*	*19*	*1828*	*218*	*648*
	JH3-6-8-IF1	1379	577	1803	109	71	21.7	29.8	0.1	27.1	0.2	1.5	3.1	2.0	9.3
		31724	*22505*	*72103*	*2619*	*3923*	*1367*	*1969*	*7*	*5600*	*26*	*113*	*432*	*175*	811
	JH3-6-8-IF3	4040	305	703	104	8	27.7	12.1	nd	28.3	nd	2.0	1.6	nd	2.3
		92916	*11879*	*28120*	*2497*	*432*	*1748*	*801*	*nd*	*5860*	*nd*	*146*	*228*	*nd*	*198*
Lwh	**JH3-1-IF2**	2201	686	826	559	51	45.1	66.6	6.6	52.4	nd	nd	7.7	nd	4.8
		50631	*26761*	*33043*	*13418*	*2821*	*2839*	*4394*	*701*	*10854*	*nd*	*nd*	*1074*	*nd*	*425*
	JH3-2-IF3	1482	439	1519	380	63	43.2	75.9	3.0	16.7	0.4	3.8	4.4	0.8	2.7
		34086	*17112*	*60749*	*9110*	*3473*	*2723*	*5007*	*322*	*3452*	*51*	*281*	*618*	*71*	*237*
	JH3-2-IF5	1633	688	1429	108	155	51.3	146.3	4.7	24.1	0.5	4.7	2.9	1.2	4.3
		37558	*26840*	*57164*	*2580*	*8502*	*3231*	*9656*	*506*	*4996*	*58*	*354*	*402*	*101*	*373*
	JH3-4-IF1	1389	562	1047	682	84	45.5	146.4	2.5	34.1	0.5	5.1	2.9	1.3	2.2
		31956	*21929*	*41878*	*16362*	*4620*	*2865*	*9664*	*267*	*7054*	*67*	*382*	*404*	*109*	*192*
	JH3-4-IF4	1034	710	1507	532	42	30.5	33.1	6.5	22.6	0.4	2.8	0.9	2.0	4.3
		23778	*27675*	*60279*	*12764*	*2319*	*1921*	*2186*	*691*	*4682*	*49*	*213*	*126*	*173*	*373*
	JH3-5-IF1	1451	394	717	744	248	48.0	352.2	8.7	23.1	4.7	3.7	26.6	1.0	6.1
		33367	*15385*	*28681*	*17853*	*13621*	*3023*	*23248*	*929*	*4785*	*578*	*279*	*3691*	*87*	*535*

Table 6. Comparison of LA-ICP-MS data obtained on brines from Jbel Haïmer and Ag-base metal brines from Roc Blanc (data from Essarraj et al., 2017a). Concentrations are in mmol/kg of solution in the first line, and ppm (second line, italic font).

FI Type	Na	Mg	K	Ca	Mn	Sr	Ag	Ba	Pb	Cu	Zn	As	Sb
Lws Jbel Haïmer	2735	73	397	941	181	6.5	0.2	10	72.2	19.1	159.1	1	0.4
	62877	*1774*	*15522*	*37713*	*9944*	*570*	*22*	*1373*	*14960*	*1214*	*10402*	*75*	*49*
Lwh Jbel Haïmer	1532	501	580	1174	107	4.1	5.3	7.6	28.8	43.9	136.8	4	1.3
	35220	*12177*	*22677*	*47052*	*5878*	*359*	*572*	*1044*	*5967*	*2790*	*8944*	*300*	*158*
Lws2 Roc Blanc	3303	470	274	819	24	33.2	0.6	4.8	14.3	4	24.1	1.5	nd
	75924	*11417*	*10730*	*32814*	*1301*	*2909*	*65*	*659*	*2963*	*254*	*1576*	*112*	*nd*
Lw Roc blanc	2047	82	196	564	18	7.7	3.5	0.9	14	2.9	5.1	2.2	9.6
	47060	*1993*	*7663*	*22604*	*989*	*675*	*378*	*124*	*2901*	*184*	*333*	*165*	*1169*

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